

EBAF Ed4.2.1 Data vs. Theory

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Budapest, Hungary

Spring CERES Science Team Meeting
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Remote Presentation

Theory: Schwarzschild (1906)

Plane-parallel, gray, LTE

Ueber das Gleichgewicht der Sonnenatmosphäre

Von

K. Schwarzschild.

Vorgelegt in der Sitzung vom 13. Januar 1906.

On the Equilibrium of the Sun's Atmosphere

by K. Schwarzschild

(Presented at the meeting of the Berlin Academy of Sciences on January 13, 1906)

$$E = \frac{A_0}{2} (1 + \bar{\tau}), \quad A = \frac{A_0}{2} (2 + \bar{\tau}), \quad B = \frac{A_0}{2} \bar{\tau}. \quad (11)$$

E emission of a layer, A upward beam, B downward beam, A_0 emerging flux at TOA, τ optical depth

$$\mathbf{Eq. (1) \quad A - E = \Delta A = A_0/2}$$

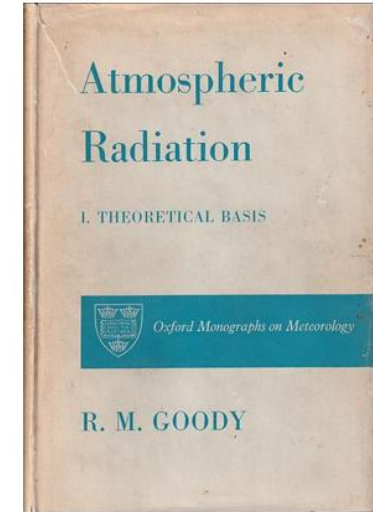
Net radiation at the surface, independent of τ

R. Goody: Atmospheric radiation. Theoretical basis. Oxford, 1964

$$F_\nu/2\pi = J_\nu(0) - B_\nu^*(0) = B_\nu^*(\tau_\nu^*) - J_\nu(\tau_\nu^*). \quad (2.115)$$

Note that if the medium is in thermodynamic equilibrium ($J_\nu = B_\nu$), equation (2.115) requires a discontinuity in B_ν (i.e. a *temperature discontinuity*) at both boundaries.

the step in B always equals $F/2\pi$



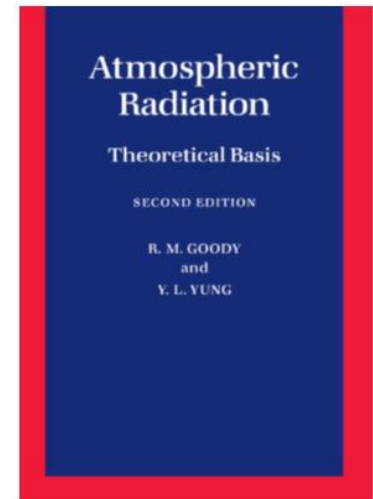
Goody – Yung: Atmospheric radiation. Theoretical basis. Oxford, 1989

$$F/2\pi = B(0) - B^*(0) = B^*(\tau_1) - B(\tau_1). \quad (2.146)$$

Equation (2.146) requires a discontinuity in the Planck function, implying a discontinuity of temperature, at the boundary.

2. There are discontinuities,

$$\Delta B = \frac{F_s}{2\pi} \quad \text{DISCONTINUITY} = \text{OLR}/2$$



John Houghton: The Physics of Atmospheres

Cambridge (1977, 1986, 2002) Eq. 2.13

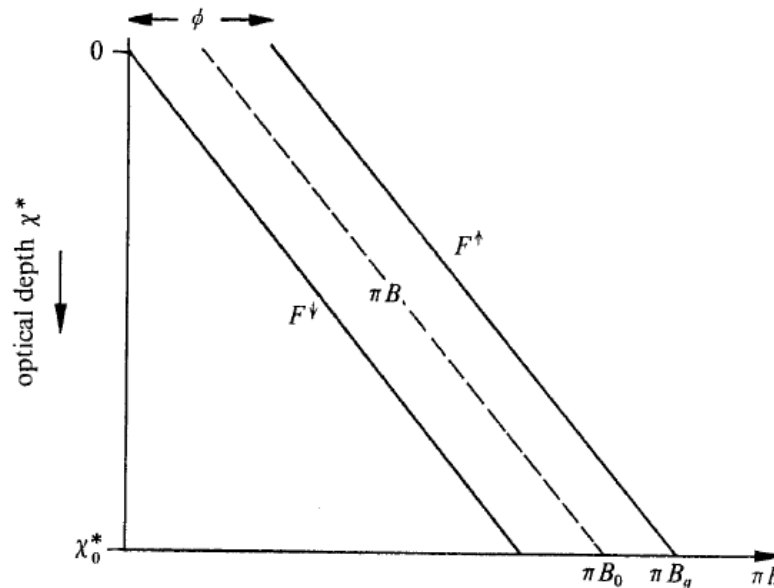
At the bottom of the atmosphere where $\chi^* = \chi_0^*$, $F^\uparrow = \pi B_g$, B_g being the black-body function at the temperature of the ground. It is easy to show that there must be a temperature discontinuity at the lower boundary, the black-body function for the air close to the ground being B_0 , and

$$B_g - B_0 = \frac{\phi}{2\pi}$$

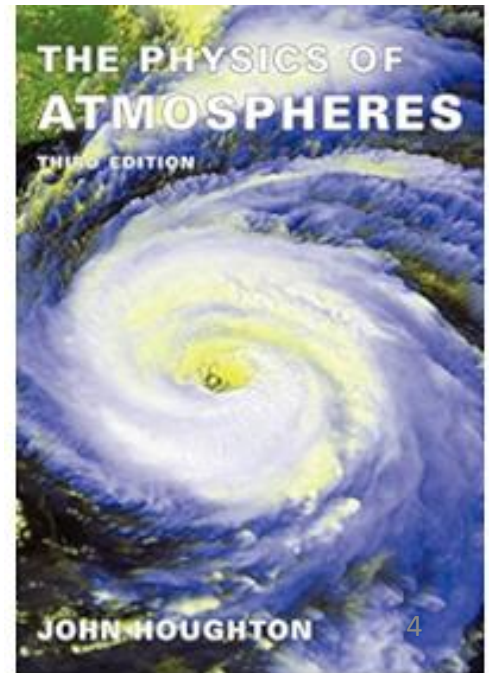
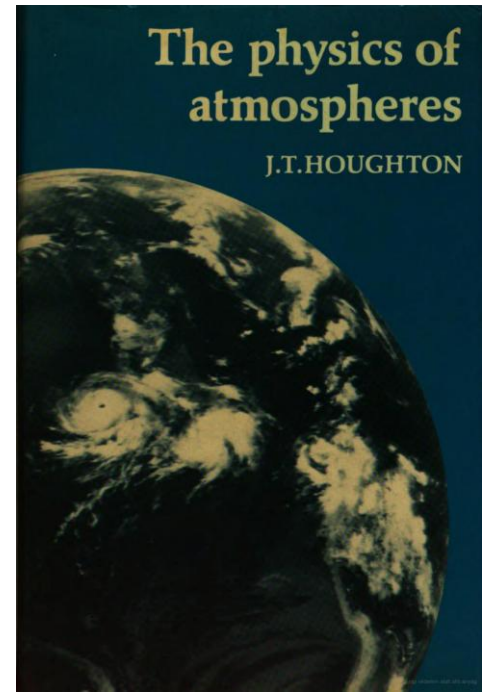
DISCONTINUITY = OLR/2
(independent of the optical depth)

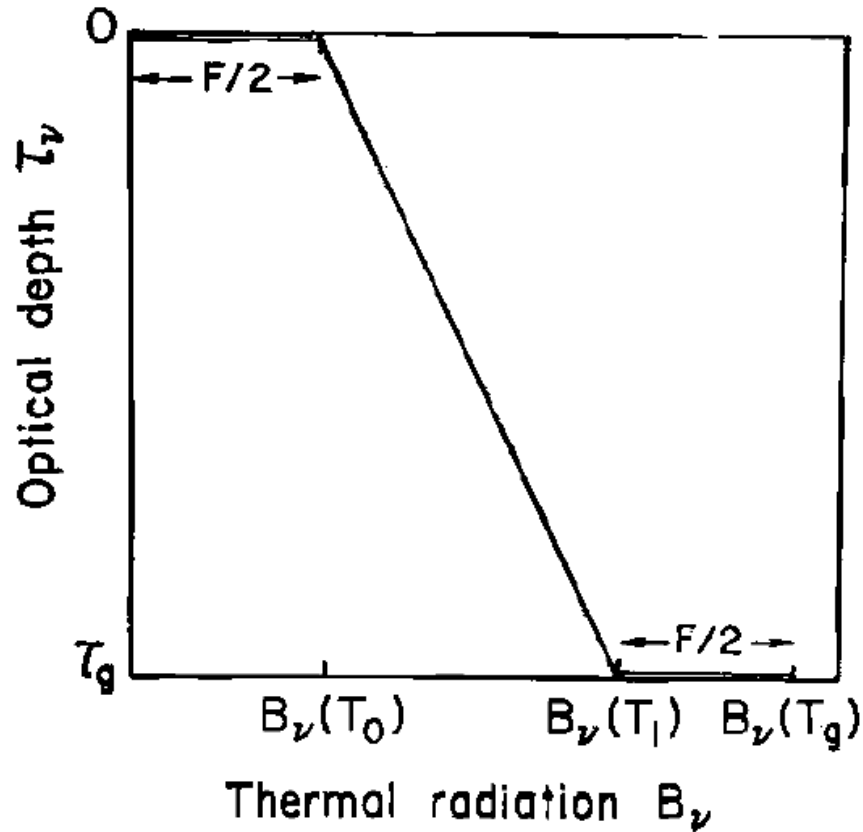
such a steep lapse rate is very unstable with respect to vertical motion, and will soon be destroyed by the process of *convection*

$$\text{Discontinuity} = \text{Convection} = \text{OLR}/2$$



(2.13)





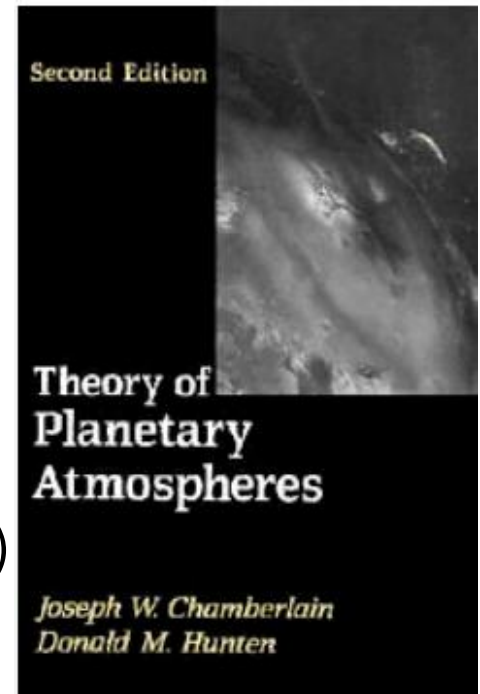
Joseph Chamberlain
 Theory of Planetary Atmospheres (1978, 1987)
 Academic Press, Fig. 1.4, Eq. 1.2.29

Fig. 1.4 The MRE solution for $T(\tau)$, presented as $B_\nu(T)$ vs. τ . Note the discontinuity at the ground and the finite skin temperature at $\tau = 0$.

**DISCONTINUITY = OLR/2
 (independent of τ)**

Hence the upward intensity at the ground is

$$I_g^+ \equiv B_\nu(T_g) = B_\nu(T_1) + \frac{1}{2}F_\nu \quad (1.2.29)$$



Visconti: Fundamentals of Physics and Chemistry of the Atmospheres
Springer (2001)

$$T^4 = \frac{T_c^4}{2} \left(\frac{3}{2} \tau + 1 \right) \quad (3.47)$$

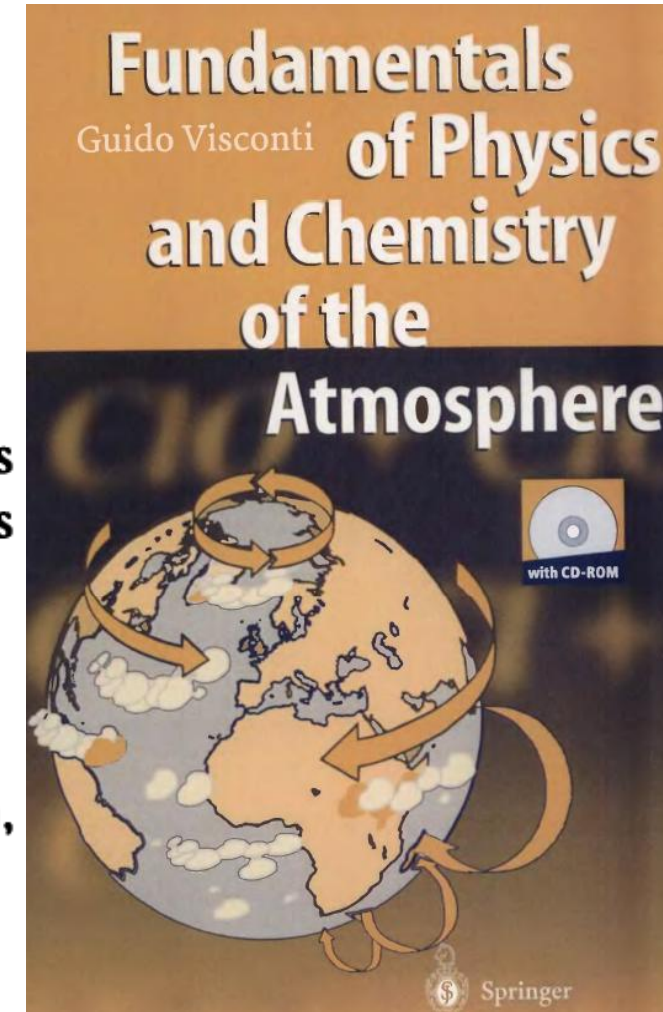
where τ has been replaced by its effective value. An interesting consequence of this solution is to discover what happens at the surface, where the optical thickness is τ^* and the temperature is T_s ; we have

$$F^\uparrow(\tau^*) = \sigma T_s^4 = \frac{1}{2} \sigma T_c^4 \left(\frac{3}{2} \tau^* + 2 \right) \quad (3.48)$$

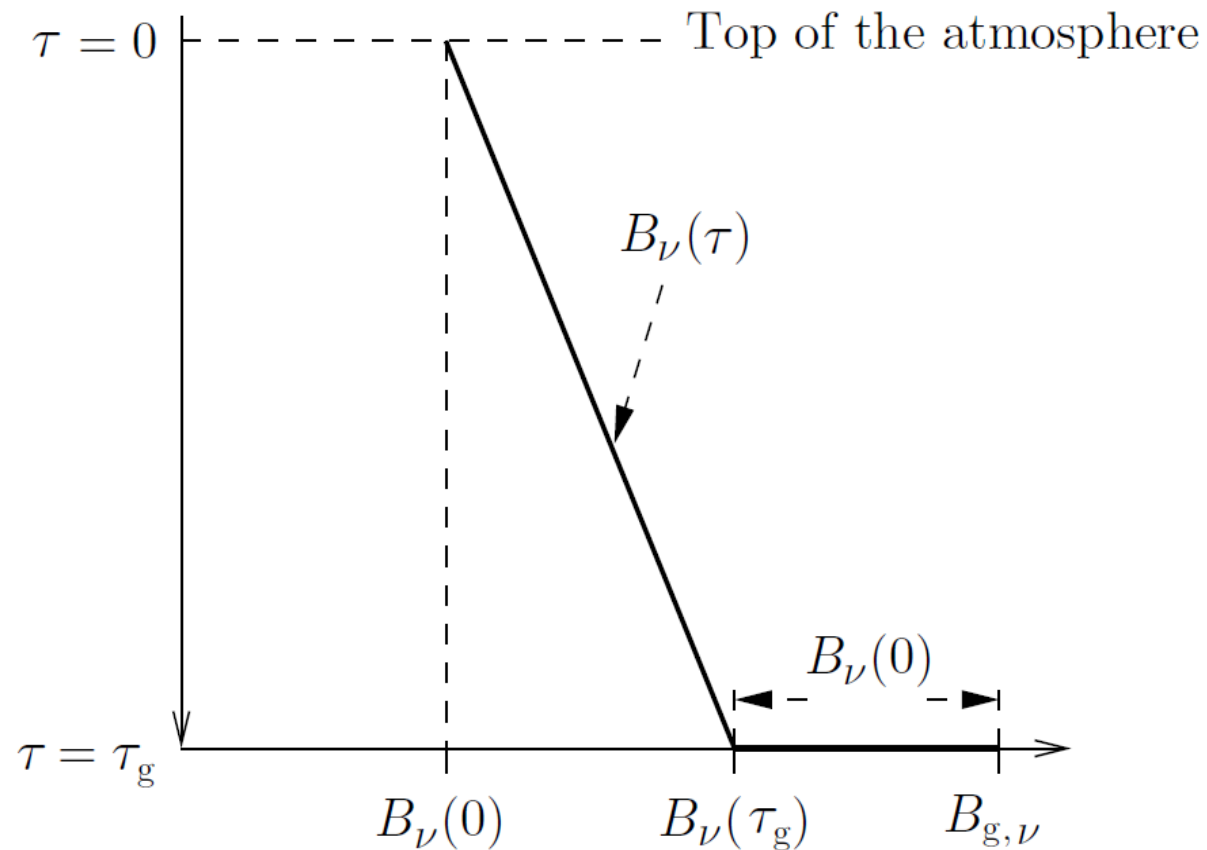
The temperature at the bottom of the atmosphere at τ^* is given by Equation (3.47), so that we have a discontinuity between the air temperature and that of the surface

$$T_s^4 - T(\tau^*) = T_c^4 / 2 \quad (3.49)$$

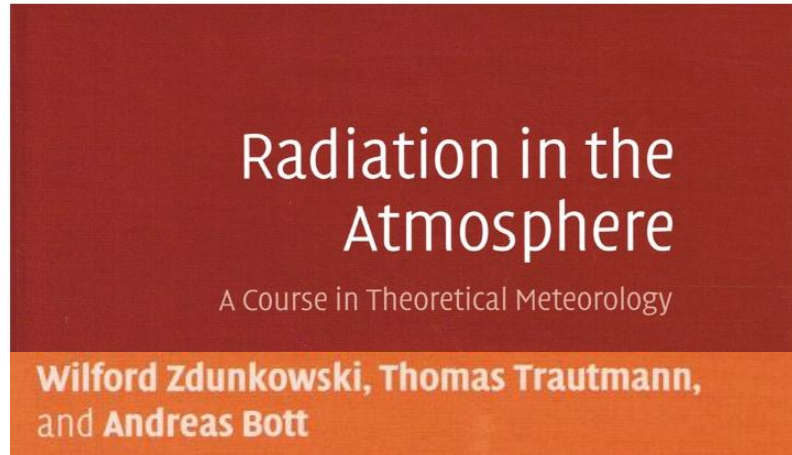
DISCONTINUITY = OLR/2



Two-stream methods for the solution of the RTE



Radiation in the Atmosphere
Zdunkowski, Trautmann and Bott
 Cambridge (2008) Fig. 6.7



DISCONTINUITY = OLR/2

$$B_{g,\nu} = B_{\nu}(\tau_g) + \frac{1}{2\pi} E_{\text{net},\nu}$$

observe a discontinuity of the curve expressing a temperature jump ΔT_g between the surface temperature T_g and the lowest atmospheric layer $T(\tau_g)$, i.e.

$$\Delta T_g = T_g - T(\tau_g) > 0 \tag{6.155}$$

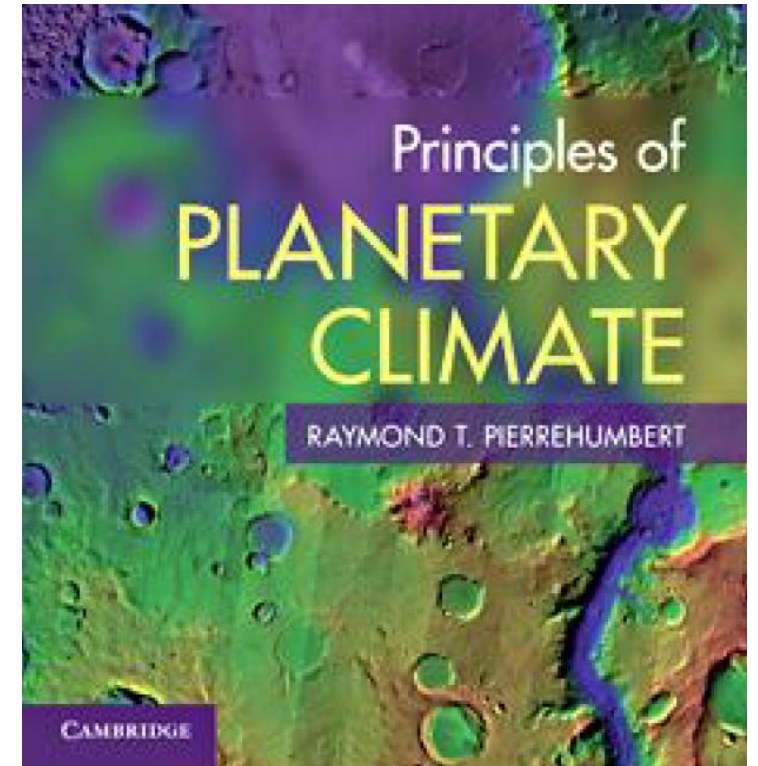
Raymond Pierrehumbert: Principles of Planetary Climate.

Cambridge (2008) Eq. 4.45

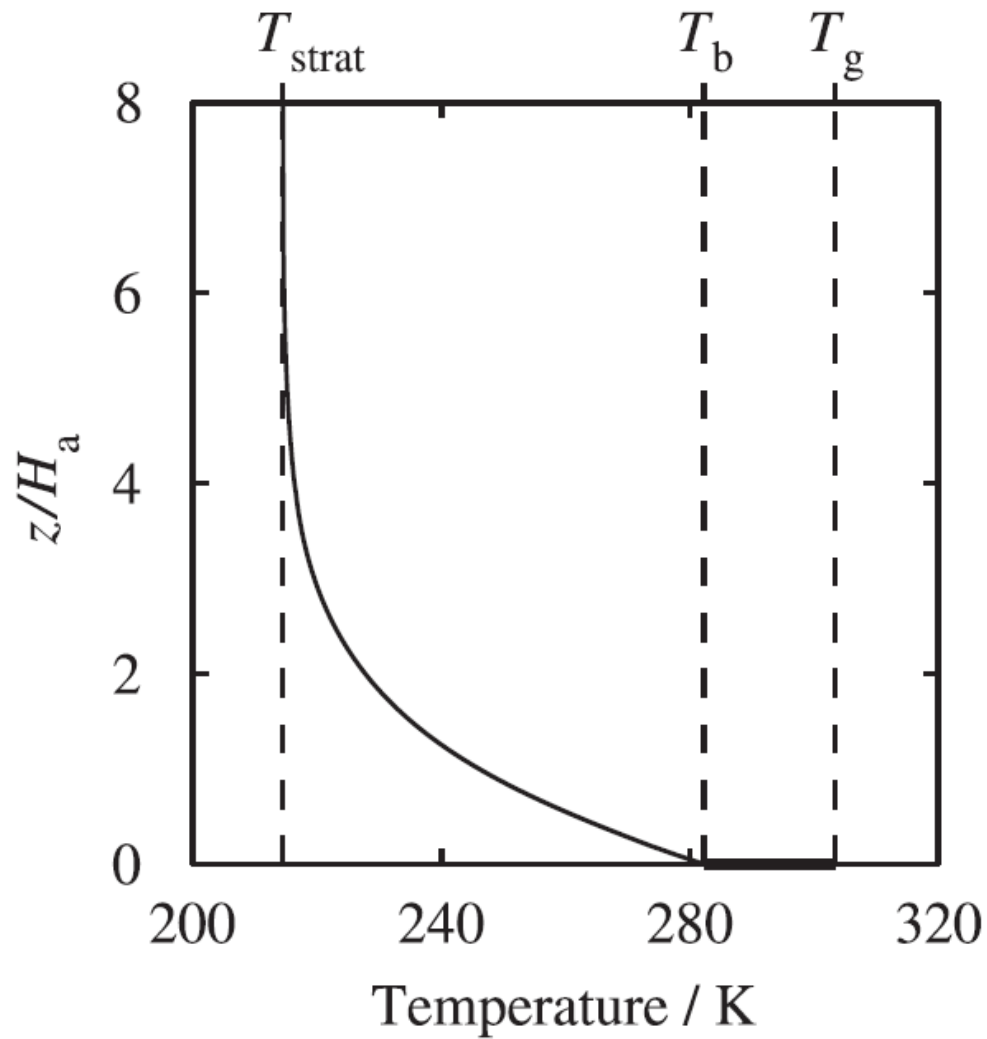
Let's now compare the surface temperature with the temperature of the air in immediate contact with the surface. From Eq. (4.42) we find that the low-level air temperature is determined by $\sigma T(0)^4 = (1 - \alpha)S \cdot \left(\frac{1}{2} + \frac{1}{2}\tau_\infty\right)$. Taking the ratio,

$$\frac{T(0)}{T_g} = \left(\frac{\frac{1}{2} + \frac{1}{2}\tau_\infty}{1 + \frac{1}{2}\tau_\infty} \right)^{1/4} \quad (4.45)$$

$$\text{DISCONTINUITY} = (1 - \alpha)S/2 = \text{OLR}/2$$



Andrews: An Introduction to Atmospheric Physics
 Cambridge (2010), pp 85-86.

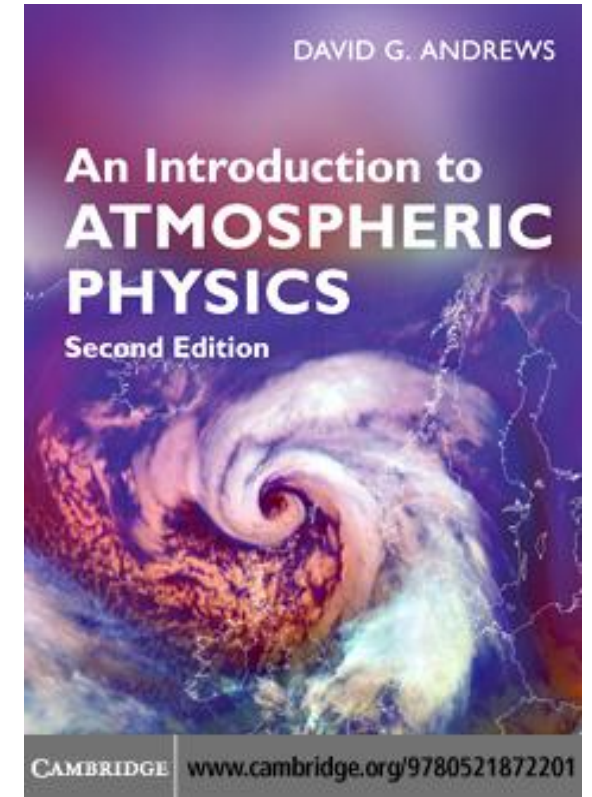


$$T_b \equiv T_e \left(\frac{1 + \chi_g^*}{2} \right)^{1/4}$$

$$T_g \equiv T_e \left(\frac{2 + \chi_g^*}{2} \right)^{1/4}$$

$$T_g^4 - T_b^4 = T_e^4/2$$

DISCONTINUITY = OLR/2



discontinuity between the bottom of the atmosphere and the ground.

Inclusion of convection in the model removes the temperature discontinuity

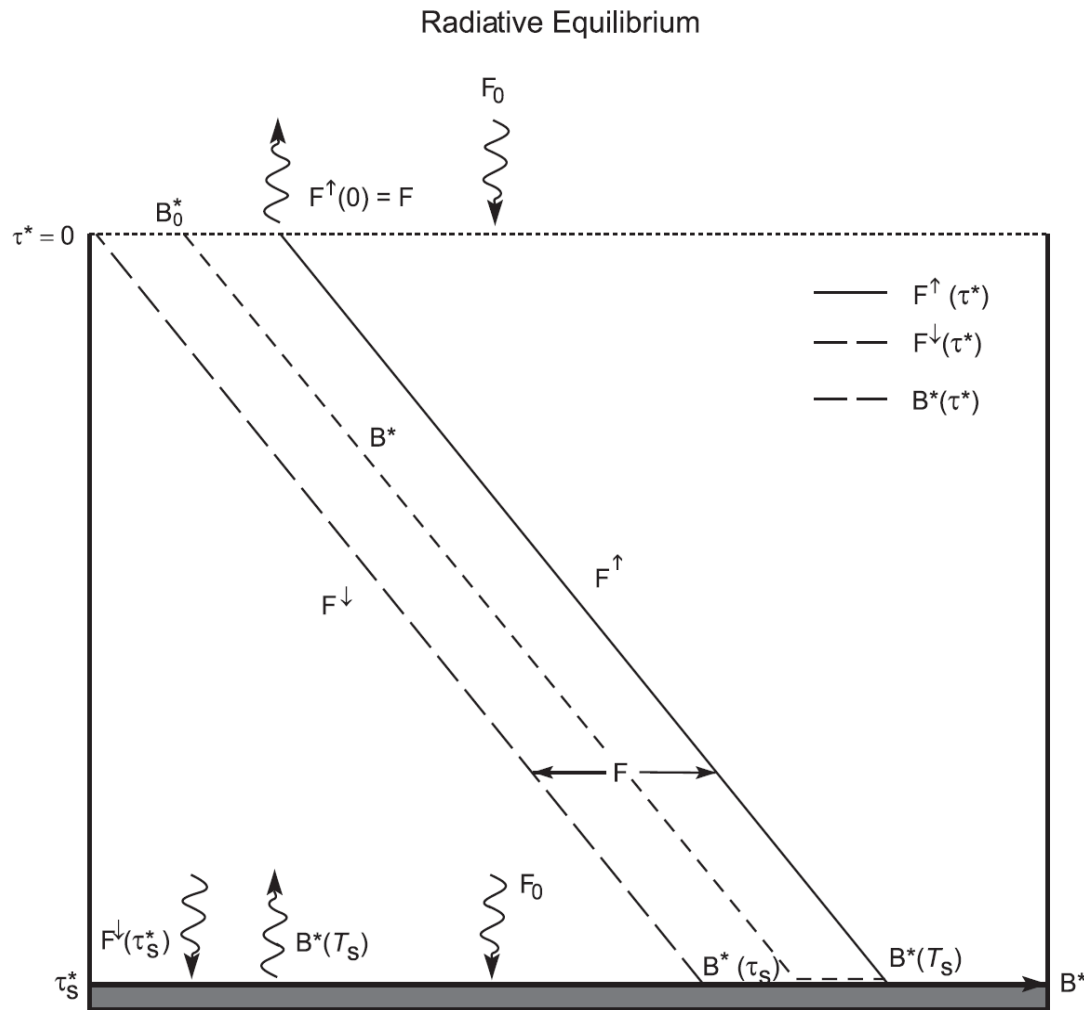
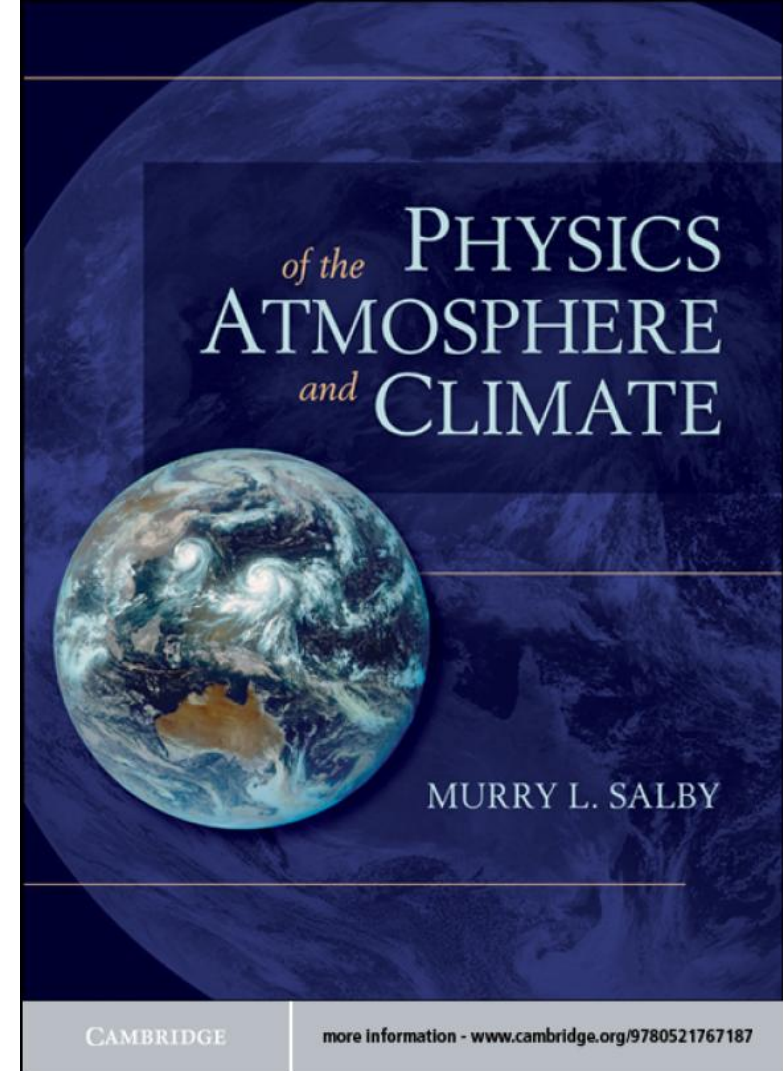


Figure 8.20 Upwelling and downwelling LW fluxes and LW emission in a gray atmosphere that is in radiative equilibrium with an incident SW flux F_0 and a black underlying surface. Note: the emission profile is discontinuous at the surface.



M. Salby: Physics of the Atmosphere and Climate
Cambridge, 2012

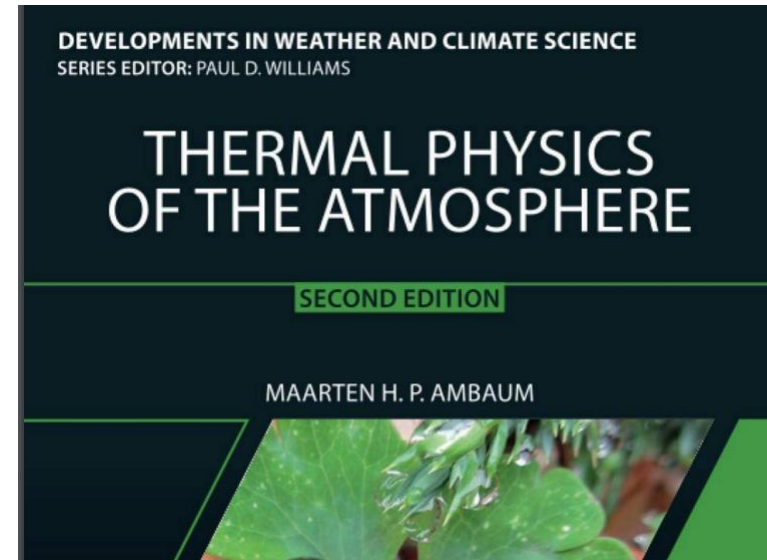
$$B^*(T_s) = B^*(\tau_s^*) + \frac{F_0}{2}. \tag{8.67}$$

According to (8.67), the temperature predicted by radiative equilibrium is discontinuous at the surface, the ground being warmer than the overlying air.



RMetS
Royal Meteorological Society

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These two budget equations can be combined by eliminating $L_{\downarrow 0}$ to find

$$\blacktriangleright \quad \sigma T_E^4 = \sigma T_S^4 + S_{n0}/2. \quad (10.56)$$

So we find that under radiative equilibrium the surface temperature is higher than the temperature of the adjoining atmosphere. This temperature discontinuity is unstable in practice and there will be turbulent heat exchange which will remove the temperature discontinuity. This instability of the radiative state is a driver of surface layer turbulence: the radiation will force the lower part of the boundary layer, the surface layer, towards instability and this tendency is compensated by turbulent fluxes near the surface.

Kevin E. Trenberth (2022): The Changing Flow of Energy Through the Climate System

Sidebar 3.2: How Does a Greenhouse Effect Work?

(page 30: Earth's Energy Balance)

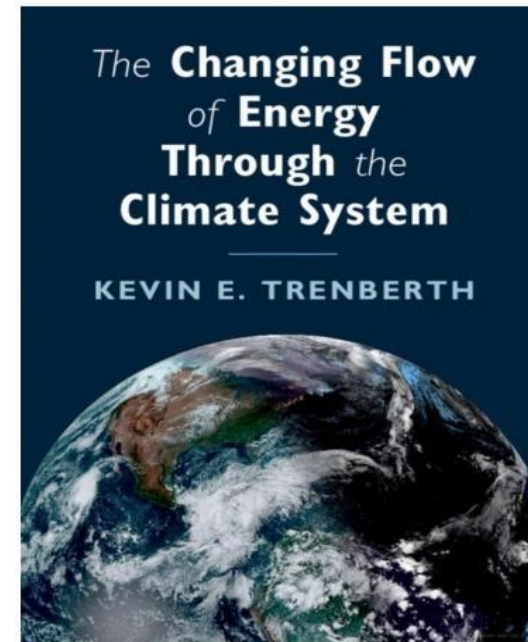
The answer is:

Imagine two plates in space

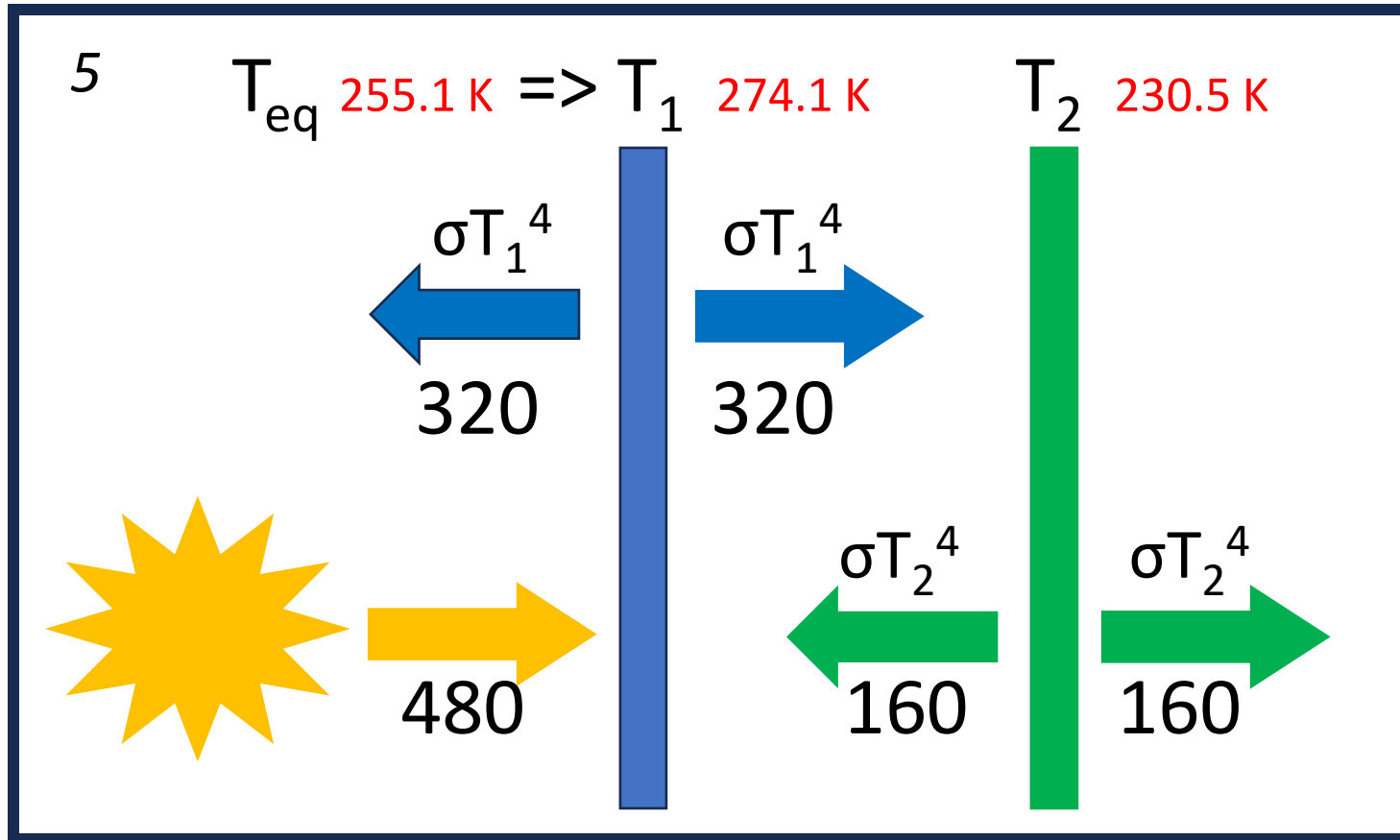
Incoming solar radiation: 240 Wm^{-2}

The plate is two-sided: 480 Wm^{-2}

Five panels, four equations.



The final state



$$G : F_2 : F_{eq} : F_1 : F_{sol} = 1 : 2 : 3 : 4 : 6$$

Trenberth's four equations

$$(T1) \quad 480 \text{ Wm}^{-2} = 2\sigma T_{eq}^4$$

$$(T2) \quad 480 \text{ Wm}^{-2} = \sigma T_1^4 + \sigma T_2^4$$

$$(T3) \quad \sigma T_1^4 = 2\sigma T_2^4$$

$$(T4) \quad 480 \text{ Wm}^{-2} = 1.5 \sigma T_1^4$$

$$F_{sol} = 480 \text{ Wm}^{-2}$$

$$F_1 = 320 \text{ Wm}^{-2}$$

$$F_{eq} = 240 \text{ Wm}^{-2}$$

$$F_2 = 160 \text{ Wm}^{-2}$$

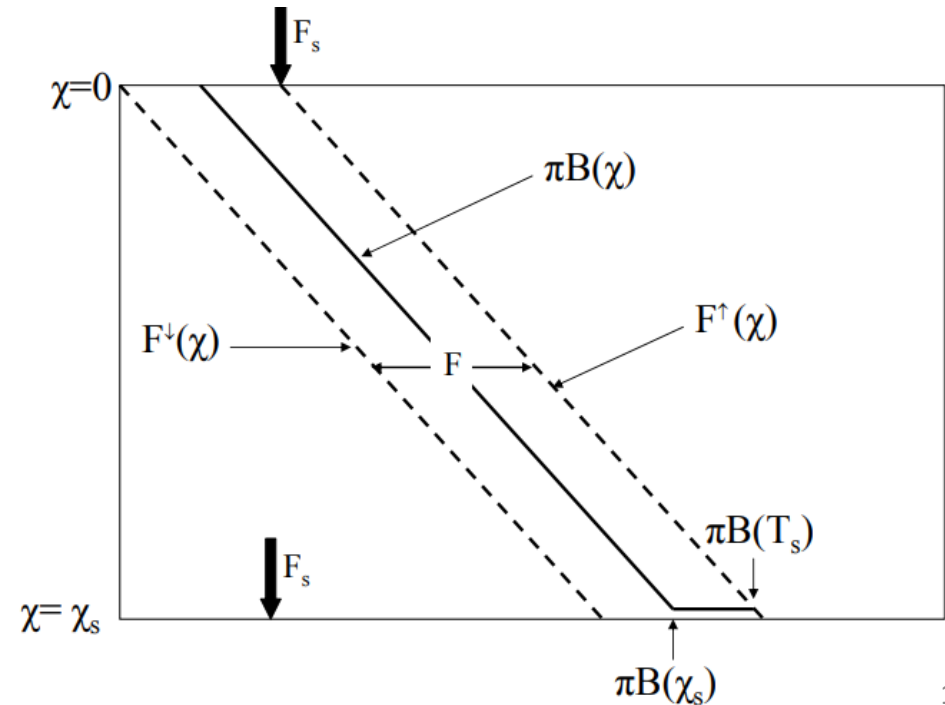
$$G = 80 \text{ Wm}^{-2}$$

University Lecture Notes

Manchester University

This expression implies that there is a temperature discontinuity between the surface and the cooler lowest layer of the atmosphere.

$$\pi B(T_s) = \pi B(\chi_s) + F_s/2.$$



Southampton University; Tyndall Centre, East Anglia

However, it is very important to notice that the (ground) surface temperature is set by F_{up} through equation (13), i.e.

$$\sigma T_g^4 = [F_{\text{up}}]_{z=0} = F_0 (1 + \tau/2) \quad (16)$$

whereas the air temperature just above the ground is set by F_{tot} through equation (11) so that

$$\sigma T_0^4 = B(\tau) = F_0 (\tau + 1)/2$$

ground surface temperature derived above exceeds that of the overlying air in this model, by an amount corresponding to an extra heat flux of $F_0/2$. This calculated ground-air temperature discontinuity may be substantial (10 or 20 °K, or more). It only occurs because we have assumed that the only heat fluxes are those due to radiation, so there is no conduction and no turbulent convection. In the real atmosphere these would operate together, as conduction will transfer heat into the air near the ground, creating an unstable

Final TP Profile

Expressions for intensity:

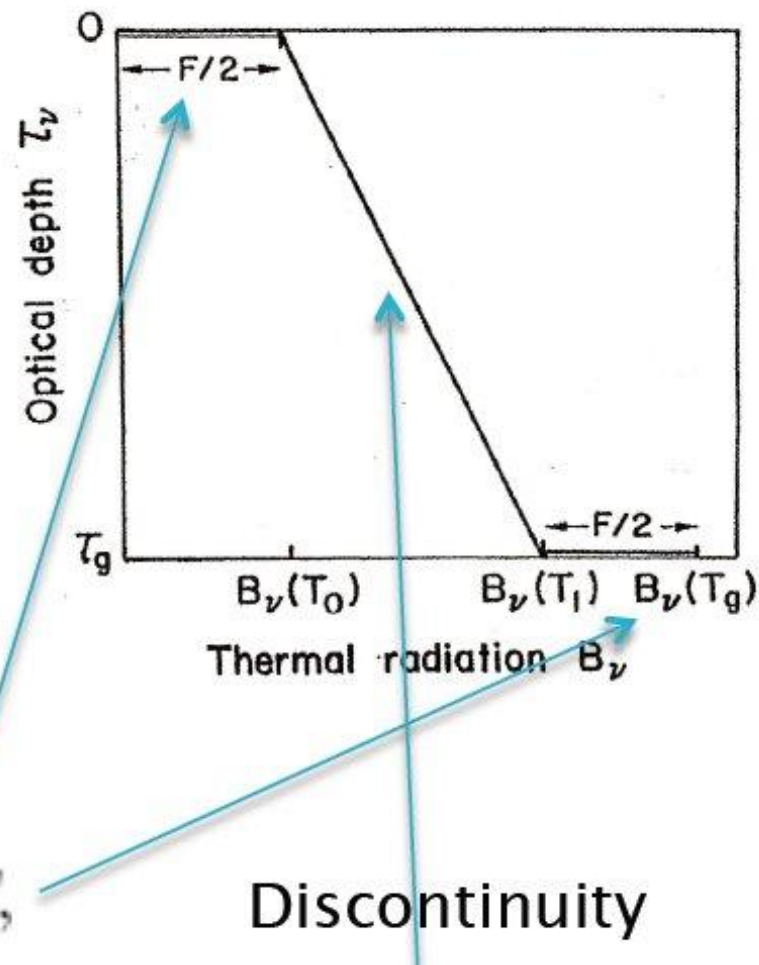
$$I^+ = B(T) + \frac{1}{2\pi}F$$

$$I^- = B(T) - \frac{1}{2\pi}F$$

Boundary Conditions:

At the Ground:

$$I_g^+ = B(T_g) = B(T_1) + \frac{1}{2\pi}F,$$



東京大学 University of Tokyo

惑星大気学_放射(2022).docx

(3.10)(3.11)より

$$\underbrace{B^*(T_S)}_{\text{地表面}} = \underbrace{B^*(\tau_s^*)}_{\text{大気下端}} + \frac{F^0}{2}$$

放射平衡では大気下端の温度と地表面温度は不連続になる。($B^* = \sigma T^4$ に注意)

From (3.10)(3.11)

$$\underbrace{B^*(T_S)}_{\text{surface}} = \underbrace{B^*(\tau_s^*)}_{\text{bottom of atmosphere}} + \frac{F^0}{2}$$

A temperature discontinuity exists at the surface. (Note that $B^* = \sigma T^4$)

University of Toronto

“We will use “Schwarzchild’s Equation for Fluxes” in our simple greenhouse atmospheric model.”

Two-Stream Model Solution - 1

- We have thus solved for the upward and downward flux densities:

$$F_{\text{up}} = \sigma T_e^4 (\chi^* + 2) / 2$$
$$F_{\text{down}} = \sigma T_e^4 \chi^* / 2$$

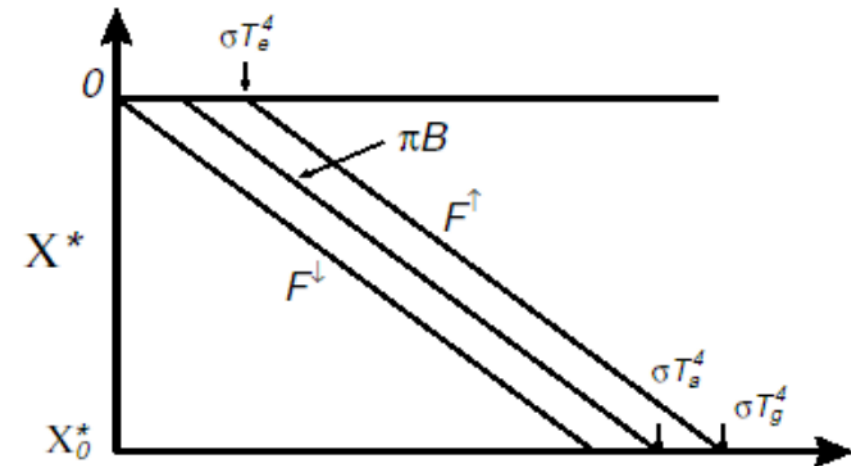
- Finally, we can use the original equations to solve for:

$$\pi B = \sigma T_a^4 = \frac{\sigma T_e^4 (\chi^* + 1)}{2}$$

where T_a is the atmospheric temperature.

- These three solutions can be plotted against χ^* .
 - Gives three parallel lines that show how the flux densities and the blackbody function vary with χ^* and hence with altitude.

Two-Stream Model Solution - 3



- If the central line represents the atmospheric temperature and the line for F_{up} must match the ground temperature, then there is a temperature discontinuity at the ground.

Harvard (2018)

We have got the temperature structure in the atmosphere as a function of τ . Now consider energy balance at the surface (looks familiar?),

$$B(T_s) = F_0 + F^\downarrow(\tau_s)$$

From their definitions, we have

$$F^\downarrow = \frac{1}{2}(\bar{F} - F)$$

As the net flux F is constant and equal to F_0 , and use Eq. (14), we have:

$$B(T_s) = B(\tau_s) + \frac{F_0}{2}$$

Note the jump at the surface.

Graeme L. Stephens: Radiative Transfer Notes AT 622.

Colorado State University, Fort Collins, CO (2003)

Example 6.3: Skin temperatures and temperature discontinuities

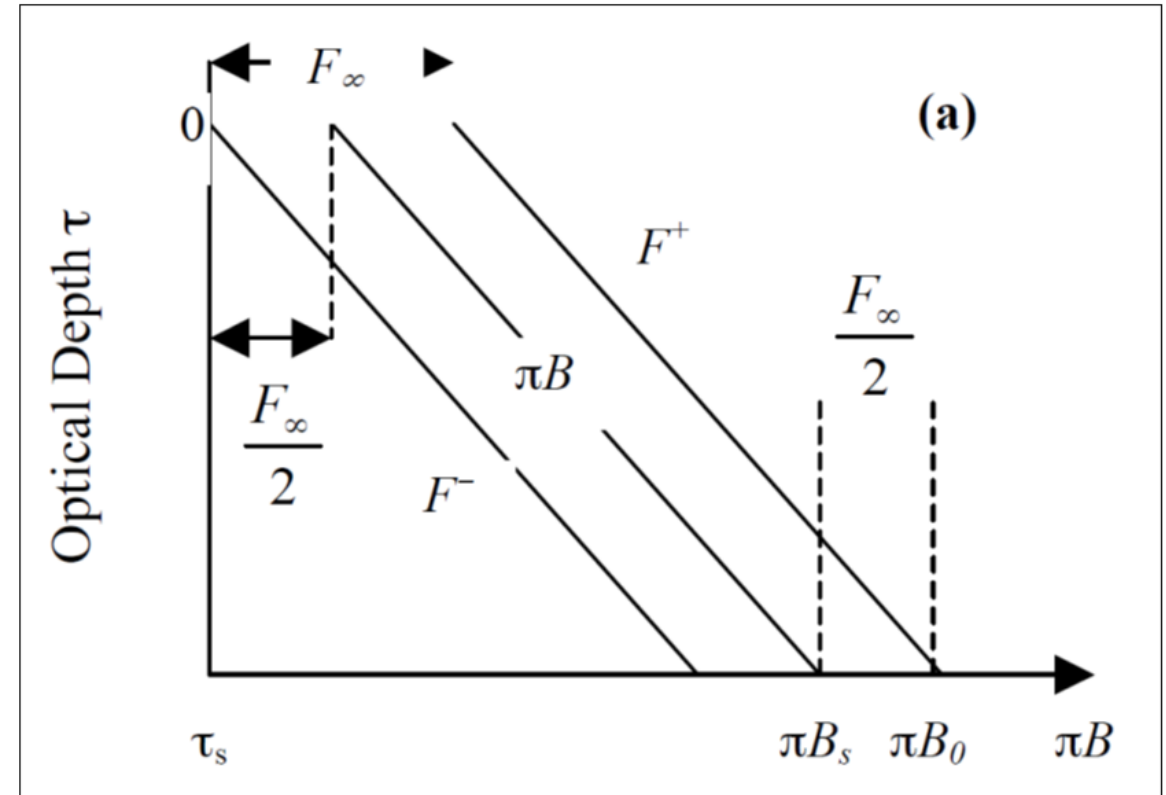
The solutions represented by Eqns. (6.10a) and (6.10b) provide rather interesting insights into the temperature profiles that are predicted by these equations. One of the results of this model is an estimate of the 'skin' temperature, which we think of as a measure of the stratospheric temperature. We obtain this using Eqn. (6.10a) with $\tilde{\tau} = 0$

$$\sigma T^4(\tilde{\tau} = 0) = \frac{F_\infty}{2}$$

and with $F_\infty \approx 235 \text{ Wm}^{-2}$, it follows that this temperature is $T_{skin} = [117.5/5.68 \times 10^{-8}]^{0.25} = 213 \text{ K}$.

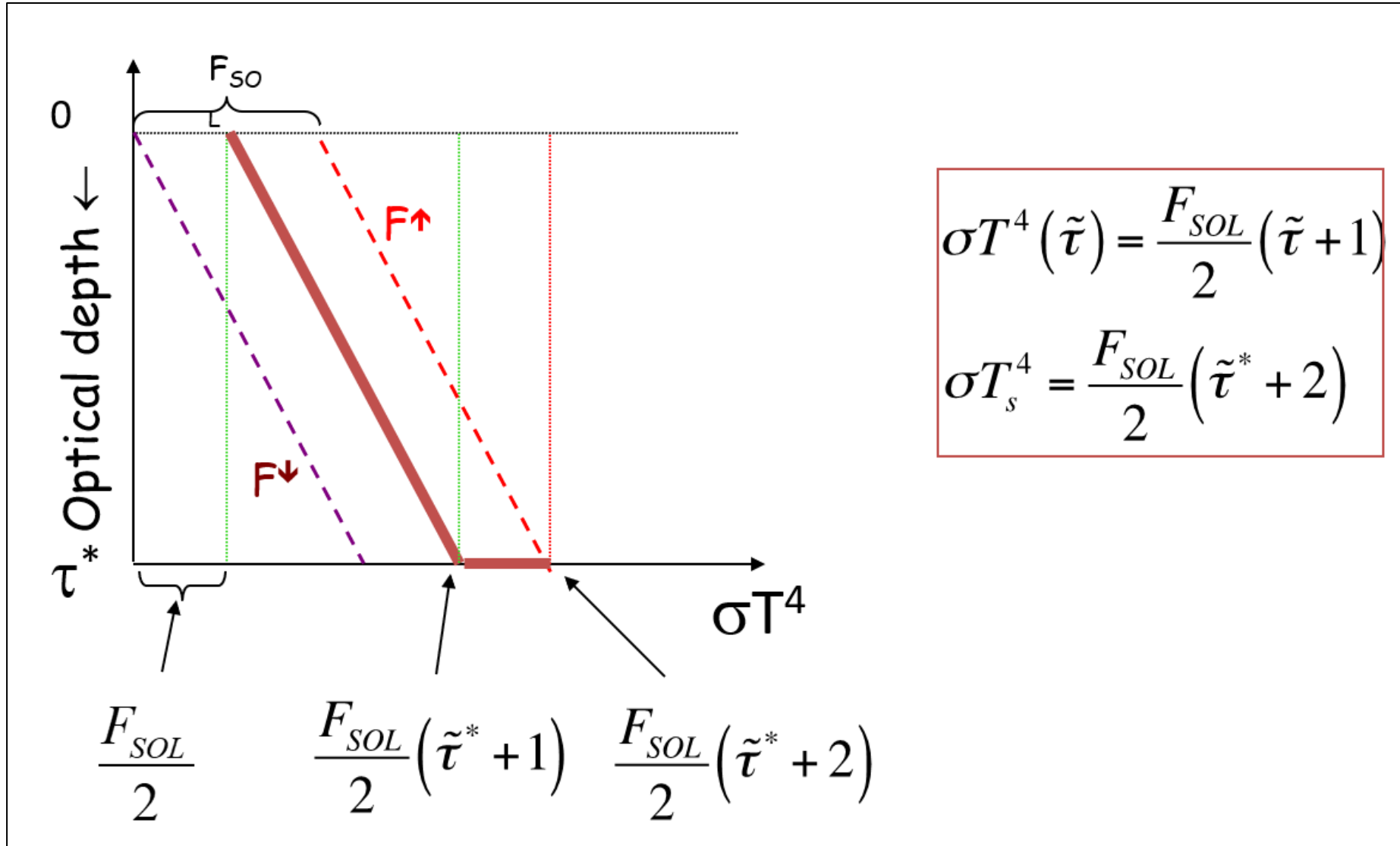
The solutions in Eqns. (6.10a) and (6.10b) predict a discontinuity between the surface temperature T_s and the air temperature just above the ground $T(\tilde{\tau}_s)$. Differencing these equations and with $\tilde{\tau} = \tilde{\tau}_s$,

$$\sigma T_s^4 - \sigma T^4(\tilde{\tau}_s) = \frac{F_\infty}{2}.$$



“This radiative equilibrium profile is unstable w.r.t. vertical motion and is destroyed by convection”

Colorado State Univ. (2013)



Early verification

CERES_EBAF-Surface_Ed2.8 Data Quality Summary (March 27, 2015)

Table 4-1. Global annual mean fluxes using data from March 2000 through February 2010 (W m⁻²).

$$\begin{array}{l} \text{Eq. (1)} \qquad \qquad \mathbf{A} \quad - \quad \mathbf{E} \qquad = \quad \mathbf{\Delta A} \qquad = \quad \mathbf{A_0/2} \\ \text{SFC} \quad \text{SW dn} - \text{SW up} + \text{LW dn} - \text{LW up} \qquad = \text{TOA LW} / 2 \\ \qquad \qquad \mathbf{(243.9 \quad - 29.7 \quad + 316.0) - 398.0} \qquad = \mathbf{132.85} \qquad = \mathbf{265.7 / 2} \end{array}$$

Early verification

CERES_EBAF-Surface_Ed2.8 Data Quality Summary (March 27, 2015)

Table 4-1. Global annual mean fluxes using data from March 2000 through February 2010 (W m^{-2}).

$$\begin{array}{l} \text{Eq. (1)} \qquad \qquad \mathbf{A} - \mathbf{E} \qquad = \Delta\mathbf{A} \qquad = \mathbf{A}_0/2 \\ \text{SFC SW dn - SW up + LW dn - LW up} \qquad = \text{TOA LW} / 2 \\ \mathbf{(243.9 - 29.7 + 316.0) - 398.0} \qquad = \mathbf{132.85} \qquad = \mathbf{265.7 / 2} \quad - \mathbf{0.65 \text{ Wm}^{-2}} \end{array}$$

Early verification

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CERES EBAF Ed2.8, 192 months (Mar 2000 – Feb 2016)

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Early verification

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CERES EBAF Ed2.8, 192 months (Mar 2000 – Feb 2016)

$$\begin{array}{l} \text{Eq. (1)} \qquad \qquad \mathbf{A} \quad - \quad \mathbf{E} \qquad = \quad \Delta\mathbf{A} \qquad = \quad \mathbf{A}_0/2 \\ \text{SFC} \quad \text{SW dn} - \text{SW up} + \text{LW dn} - \text{LW up} \qquad = \text{TOA LW} / 2 \\ \qquad \qquad \qquad \mathbf{(244.06 - 29.74 + 316.27) - 398.40} \qquad = \mathbf{132.19} \qquad = \mathbf{265.59 / 2} \quad - \mathbf{0.60 \text{ Wm}^{-2}} \end{array}$$

Early verification

CERES_EBAF-Surface_Ed2.8 Data Quality Summary (March 27, 2015)

Table 4-1. Global annual mean fluxes using data from March 2000 through February 2010 (W m^{-2}).

$$\begin{array}{l} \text{Eq. (1)} \qquad \qquad \mathbf{A} \quad - \quad \mathbf{E} \qquad = \quad \Delta\mathbf{A} \qquad = \quad \mathbf{A}_0/2 \\ \text{SFC} \quad \text{SW dn} - \text{SW up} + \text{LW dn} - \text{LW up} \qquad = \text{TOA LW} / 2 \\ \qquad \qquad \qquad \mathbf{(243.9 \quad - 29.7 \quad + 316.0) - 398.0} \qquad = \mathbf{132.85} \qquad = \mathbf{265.7 / 2} \quad - \mathbf{0.65 \text{ Wm}^{-2}} \end{array}$$

CERES EBAF Ed2.8, 192 months (Mar 2000 – Feb 2016)

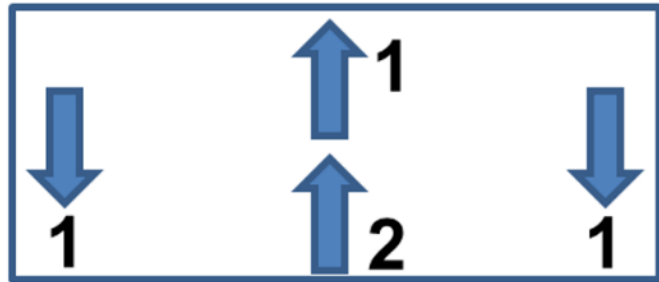
$$\begin{array}{l} \text{Eq. (1)} \qquad \qquad \mathbf{A} \quad - \quad \mathbf{E} \qquad = \quad \Delta\mathbf{A} \qquad = \quad \mathbf{A}_0/2 \\ \text{SFC} \quad \text{SW dn} - \text{SW up} + \text{LW dn} - \text{LW up} \qquad = \text{TOA LW} / 2 \\ \qquad \qquad \qquad \mathbf{(244.06 - 29.74 + 316.27) - 398.40} \qquad = \mathbf{132.19} \qquad = \mathbf{265.59 / 2} \quad - \mathbf{0.60 \text{ Wm}^{-2}} \end{array}$$

Earth heat uptake for July 2005–June 2010 is $0.58 \pm 0.38 \text{ Wm}^{-2}$

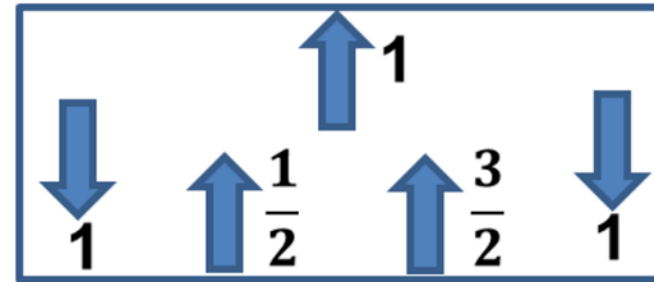
Eq. (2): Schwarzschild (1906, Eq. 11) at $\bar{\tau} = 2$

$$E = \frac{A_0}{2} (1 + \bar{\tau}), \quad A = \frac{A_0}{2} (2 + \bar{\tau}), \quad B = \frac{A_0}{2} \bar{\tau}. \quad (11)$$

Eq. (2) $A = 2A_0$



$E = 3A_0/2; \quad B = A_0$

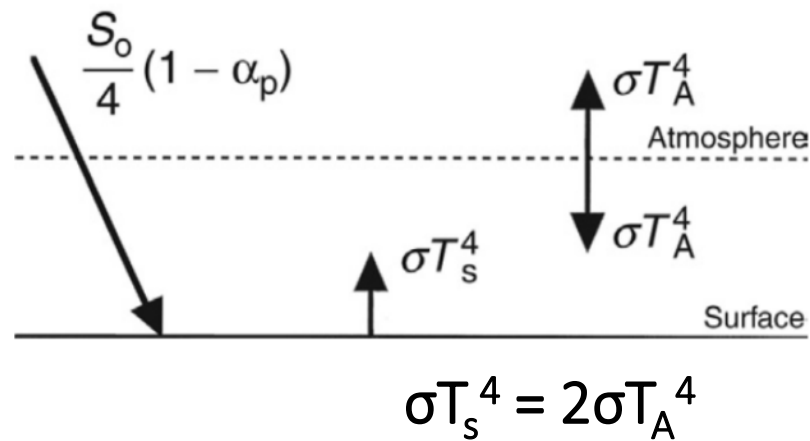


Eq. (2) $A = 2A_0$

Clear-sky

Eq.(1) $\Delta A = A - E = A_0/2$

3.9 Clouds and Radiation



Hartmann (1994) Fig. 2.3

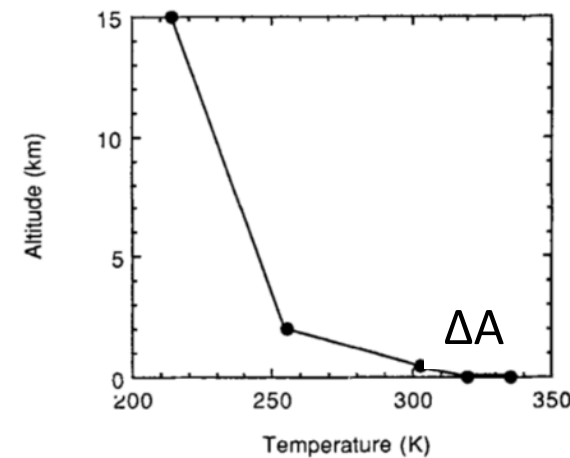


Fig. 3.11

CERES EBAF Ed2.8, Global means (Mar 2000 – Feb 2016)

Clear-sky

$$\text{Eq. (1)} \quad \begin{array}{rclclcl} \mathbf{A} & - & \mathbf{E} & = & \Delta\mathbf{A} & = & \mathbf{A}_0/2 \\ \mathbf{530.59} & - & \mathbf{398.40} & = & \mathbf{132.19} & = & \mathbf{265.59 / 2} & - \mathbf{0.60 \text{ Wm}^{-2}} \end{array}$$

$$\text{Eq. (2)} \quad \begin{array}{rcl} \mathbf{A} & = & \mathbf{2A_0} \\ \mathbf{530.59} & = & \mathbf{2 \times 265.59} \end{array}$$

CERES EBAF Ed2.8, Global means (Mar 2000 – Feb 2016)

Clear-sky

$$\text{Eq. (1)} \quad \begin{array}{rclclcl} \mathbf{A} & - & \mathbf{E} & = & \Delta\mathbf{A} & = & \mathbf{A}_0/2 \\ \mathbf{530.59} & - & \mathbf{398.40} & = & \mathbf{132.19} & = & \mathbf{265.59 / 2} & - & \mathbf{0.60 \text{ Wm}^{-2}} \end{array}$$

$$\text{Eq. (2)} \quad \begin{array}{rclcl} \mathbf{A} & = & \mathbf{2A_0} \\ \mathbf{530.59} & = & \mathbf{2 \times 265.59} & - & \mathbf{0.59 \text{ Wm}^{-2}} \end{array}$$

CERES EBAF Ed2.8, Global means (Mar 2000 – Feb 2016)

Clear-sky

$$\text{Eq. (1)} \quad \begin{array}{rclcl} \mathbf{A} - \mathbf{E} & = & \Delta\mathbf{A} & = & \mathbf{A}_0/2 \\ 530.59 - 398.40 & = & 132.19 & = & 265.59 / 2 \end{array} \quad - 0.60 \text{ Wm}^{-2}$$

$$\text{Eq. (2)} \quad \begin{array}{rcl} \mathbf{A} & = & 2\mathbf{A}_0 \\ 530.59 & = & 2 \times 265.59 \end{array} \quad - 0.59 \text{ Wm}^{-2}$$

$\Delta\mathbf{A} : \mathbf{A}_0 : \mathbf{E} : \mathbf{A} = \mathbf{1} : \mathbf{2} : \mathbf{3} : \mathbf{4}$ (clear-sky) justified within EEI

CERES EBAF Ed2.8, Global means (Mar 2000 – Feb 2016)

Clear-sky

$$\text{Eq. (1)} \quad \begin{array}{rclcl} \mathbf{A} - \mathbf{E} & = & \Delta\mathbf{A} & = & \mathbf{A}_0/2 \\ \mathbf{530.59} - \mathbf{398.40} & = & \mathbf{132.19} & = & \mathbf{265.59 / 2} \quad \mathbf{- 0.60 \text{ Wm}^{-2}} \end{array}$$

$$\text{Eq. (2)} \quad \begin{array}{rclcl} \mathbf{A} & = & 2\mathbf{A}_0 \\ \mathbf{530.59} & = & \mathbf{2 \times 265.59} \quad \mathbf{- 0.59 \text{ Wm}^{-2}} \end{array}$$

$\Delta\mathbf{A} : \mathbf{A}_0 : \mathbf{E} : \mathbf{A} = \mathbf{1} : \mathbf{2} : \mathbf{3} : \mathbf{4}$ (clear-sky) justified within EEI

$$\mathbf{E} = (3/2) \mathbf{A}_0$$

$$\mathbf{g} = 1/3$$

CERES EBAF Ed2.8, Global means (Mar 2000 – Feb 2016)

Clear-sky

$$\text{Eq. (1)} \quad \begin{array}{rclcl} \mathbf{A} & - & \mathbf{E} & = & \mathbf{\Delta A} & = & \mathbf{A_0/2} \\ \mathbf{530.59} & - & \mathbf{398.40} & = & \mathbf{132.19} & = & \mathbf{265.59 / 2} & - & \mathbf{0.60 \text{ Wm}^{-2}} \end{array}$$

$$\text{Eq. (2)} \quad \begin{array}{rclcl} \mathbf{A} & & & = & \mathbf{2A_0} \\ \mathbf{530.59} & & & = & \mathbf{2 \times 265.59} & - & \mathbf{0.59 \text{ Wm}^{-2}} \end{array}$$

$\Delta A : A_0 : E : A = \mathbf{1 : 2 : 3 : 4}$ (clear-sky) justified within EEI

$$\begin{array}{l} \mathbf{E = (3/2) A_0} \quad \mathbf{398.40 = (3/2) \times 265.59/2} \\ \mathbf{g = 1/3} \quad \quad \quad \mathbf{(398.40 - 265.59) / 398.40} \end{array}$$

CERES EBAF Ed2.8, Global means (Mar 2000 – Feb 2016)

Clear-sky

$$\text{Eq. (1)} \quad \begin{array}{rclclcl} \mathbf{A} & - & \mathbf{E} & = & \mathbf{\Delta A} & = & \mathbf{A_0/2} \\ \mathbf{530.59} & - & \mathbf{398.40} & = & \mathbf{132.19} & = & \mathbf{265.59 / 2} & - & \mathbf{0.60 \text{ Wm}^{-2}} \end{array}$$

$$\text{Eq. (2)} \quad \begin{array}{rclcl} \mathbf{A} & = & \mathbf{2A_0} \\ \mathbf{530.59} & = & \mathbf{2 \times 265.59} & - & \mathbf{0.59 \text{ Wm}^{-2}} \end{array}$$

$\Delta A : A_0 : E : A = \mathbf{1 : 2 : 3 : 4}$ (clear-sky) justified within EEI

$$\begin{array}{l} \mathbf{E = (3/2) A_0} \quad \mathbf{398.40 = (3/2) \times 265.59/2 + 0.015} \\ \mathbf{g = 1/3} \quad \quad \quad \mathbf{(398.40 - 265.59) / 398.40 = 0.33336} \end{array}$$

Creating the all-sky versions

Eq. (1) **SFC Net** = $A - E = A_0/2$ (clear-sky, net)

Eq. (2) **SFC Tot** = $A = 2A_0$ (clear-sky, total at $\tau = 2$)

Separating atmospheric radiation transfer from the longwave cloud effect (LWCRE):

Eq. (3) **SFC Net** = $A - E = (A_0 - L)/2$ (all-sky, net, incl LWCRE)

Eq. (4) **SFC Tot** = $A = 2A_0 + L$ (all-sky, total at $\tau = 2$ incl LWCRE)

Verification of the four equations

CERES EBAF Ed4.1 Version 3, 22 years (April 2000 – March 2022) (Wm^{-2})

$$\begin{aligned} \text{Eq. (1)} \quad \text{SFC SW down} - \text{SW up} + \text{LW down} - \text{LW up (clear)} &= \text{TOA LW (clear)}/2 \\ 240.8680 - 29.0724 + 317.4049 - 398.5211 &= 266.0122 / 2 \end{aligned}$$

$$\begin{aligned} \text{Eq. (2)} \quad \text{SFC SW down} - \text{SW up} + \text{LW down} \quad (\text{clear}) &= 2 \times \text{TOA LW (clear)} \\ 240.8680 - 29.0724 + 317.4049 &= 2 \times 266.0122 \end{aligned}$$

$$\begin{aligned} \text{Eq. (3)} \quad \text{SFC SW down} - \text{SW up} + \text{LW down} - \text{LW up} \quad (\text{all}) &= [\text{TOA LW (all)} - \text{LWCRE}]/2 \\ 186.8544 - 23.1629 + 345.0108 - 398.7550 &= (240.2450 - 25.7672)/2 \end{aligned}$$

$$\begin{aligned} \text{Eq. (4)} \quad \text{SFC SW down} - \text{SW up} + \text{LW down} \quad (\text{all}) &= 2 \times \text{TOA LW (all)} + \text{LWCRE} \\ 186.8544 - 23.1629 + 345.0108 &= 2 \times 240.2450 + 25.7672 \end{aligned}$$

Verification of the four equations

CERES EBAF Ed4.1 Version 3, 22 years (April 2000 – March 2022) (Wm^{-2})

$$\begin{aligned} \text{Eq. (1)} \quad \text{SFC SW down} - \text{SW up} + \text{LW down} - \text{LW up (clear)} &= \text{TOA LW (clear)}/2 \\ 240.8680 - 29.0724 + 317.4049 - 398.5211 &= 266.0122 / 2 \quad - 2.3267 \end{aligned}$$

$$\begin{aligned} \text{Eq. (2)} \quad \text{SFC SW down} - \text{SW up} + \text{LW down} \quad (\text{clear}) &= 2 \times \text{TOA LW (clear)} \\ 240.8680 - 29.0724 + 317.4049 &= 2 \times 266.0122 \quad - 2.8238 \end{aligned}$$

$$\begin{aligned} \text{Eq. (3)} \quad \text{SFC SW down} - \text{SW up} + \text{LW down} - \text{LW up} \quad (\text{all}) &= [\text{TOA LW (all)} - \text{LWCRE}]/2 \\ 186.8544 - 23.1629 + 345.0108 - 398.7550 &= (240.2450 - 25.7672)/2 \quad + 2.7083 \end{aligned}$$

$$\begin{aligned} \text{Eq. (4)} \quad \text{SFC SW down} - \text{SW up} + \text{LW down} \quad (\text{all}) &= 2 \times \text{TOA LW (all)} + \text{LWCRE} \\ 186.8544 - 23.1629 + 345.0108 &= 2 \times 240.2450 + 25.7672 \quad + 2.4450 \end{aligned}$$

Verification of the four equations

CERES EBAF Ed4.1 Version 3, 22 years (April 2000 – March 2022) (Wm^{-2})

Eq. (1)	SFC SW down – SW up + LW down – LW up (clear)	= TOA LW (clear)/2		
	240.8680 – 29.0724 + 317.4049 – 398.5211	= 266.0122 /2	– 2.3267	} –2.5753
Eq. (2)	SFC SW down – SW up + LW down (clear)	= 2 × TOA LW (clear)		
	240.8680 – 29.0724 + 317.4049	= 2 × 266.0122	– 2.8238	
Eq. (3)	SFC SW down – SW up + LW down – LW up (all)	= [TOA LW (all) – LWCRE]/2		
	186.8544 – 23.1629 + 345.0108 – 398.7550	= (240.2450 – 25.7672)/2	+ 2.7083	} +2.5767
Eq. (4)	SFC SW down – SW up + LW down (all)	= 2 × TOA LW (all) + LWCRE		
	186.8544 – 23.1629 + 345.0108	= 2 × 240.2450 + 25.7672	+ 2.4450	

Verification of the four equations

CERES EBAF Ed4.1 Version 3, 22 years (April 2000 – March 2022) (Wm^{-2})

Eq. (1)	SFC SW down – SW up + LW down – LW up (clear)	= TOA LW (clear)/2		
	240.8680 – 29.0724 + 317.4049 – 398.5211	= 266.0122 /2	– 2.3267	} –2.5753
Eq. (2)	SFC SW down – SW up + LW down (clear)	= 2 × TOA LW (clear)		
	240.8680 – 29.0724 + 317.4049	= 2 × 266.0122	– 2.8238	
Eq. (3)	SFC SW down – SW up + LW down – LW up (all)	= [TOA LW (all) – LWCRE]/2		
	186.8544 – 23.1629 + 345.0108 – 398.7550	= (240.2450 – 25.7672)/2 + 2.7083		} +2.5767
Eq. (4)	SFC SW down – SW up + LW down (all)	= 2 × TOA LW (all) + LWCRE		
	186.8544 – 23.1629 + 345.0108	= 2 × 240.2450 + 25.7672 + 2.4450		
			Mean	0.0007

CERES EBAF Ed4.1 Version 3, 22 years (April 2000 – March 2022) (Wm⁻²)

CERES EBAF Ed4.2 Version 4, 22 years (April 2000 – March 2022) (Wm⁻²)

CERES EBAF Ed4.2 Version 4, 24 years (April 2000 – March 2024) (Wm⁻²)

Eq. (1)	SFC SW down – SW up + LW down – LW up (clear) = TOA LW (clear)/2	
	240.8680 – 29.0724 + 317.4049 – 398.5211 = 266.0122 /2	– 2.3267
	241.0969 – 29.7521 + 317.8744 – 398.5890 = 265.9594 /2	– 2.3495
	241.0514 – 29.7043 + 318.0984 – 398.7742 = 265.9748 /2	– 2.3161
Eq. (2)	SFC SW down – SW up + LW down (clear) = 2 × TOA LW (clear)	
	240.8680 – 29.0724 + 317.4049 = 2 × 266.0122	– 2.8238
	241.0969 – 29.7521 + 317.8744 = 2 × 265.9594	– 2.6996
	241.0514 – 29.7043 + 318.0984 = 2 × 265.9748	– 2.5042
Eq. (3)	SFC SW down – SW up + LW down – LW up (all) = [TOA LW (all) – LWCRE]/2	
	186.8544 – 23.1629 + 345.0108 – 398.7550 = (240.2450 – 25.7672)/2	+ 2.7083
	187.1451 – 23.4950 + 346.1057 – 398.4220 = (240.3317 – 25.6277)/2	+ 3.9818
	187.1756 – 23.4607 + 346.3158 – 398.6162 = (240.3894 – 25.5854)/2	+ 4.0126
Eq. (4)	SFC SW down – SW up + LW down (all) = 2 × TOA LW (all) + LWCRE	
	186.8544 – 23.1629 + 345.0108 = 2 × 240.2450 + 25.7672	+ 2.4450
	187.1451 – 23.4950 + 346.1057 = 2 × 240.3317 + 25.6277	+ 3.4647
	187.1756 – 23.4607 + 346.3158 = 2 × 240.3894 + 25.5854	+ 3.6665

Mean 0.0007

0.5994

0.7147

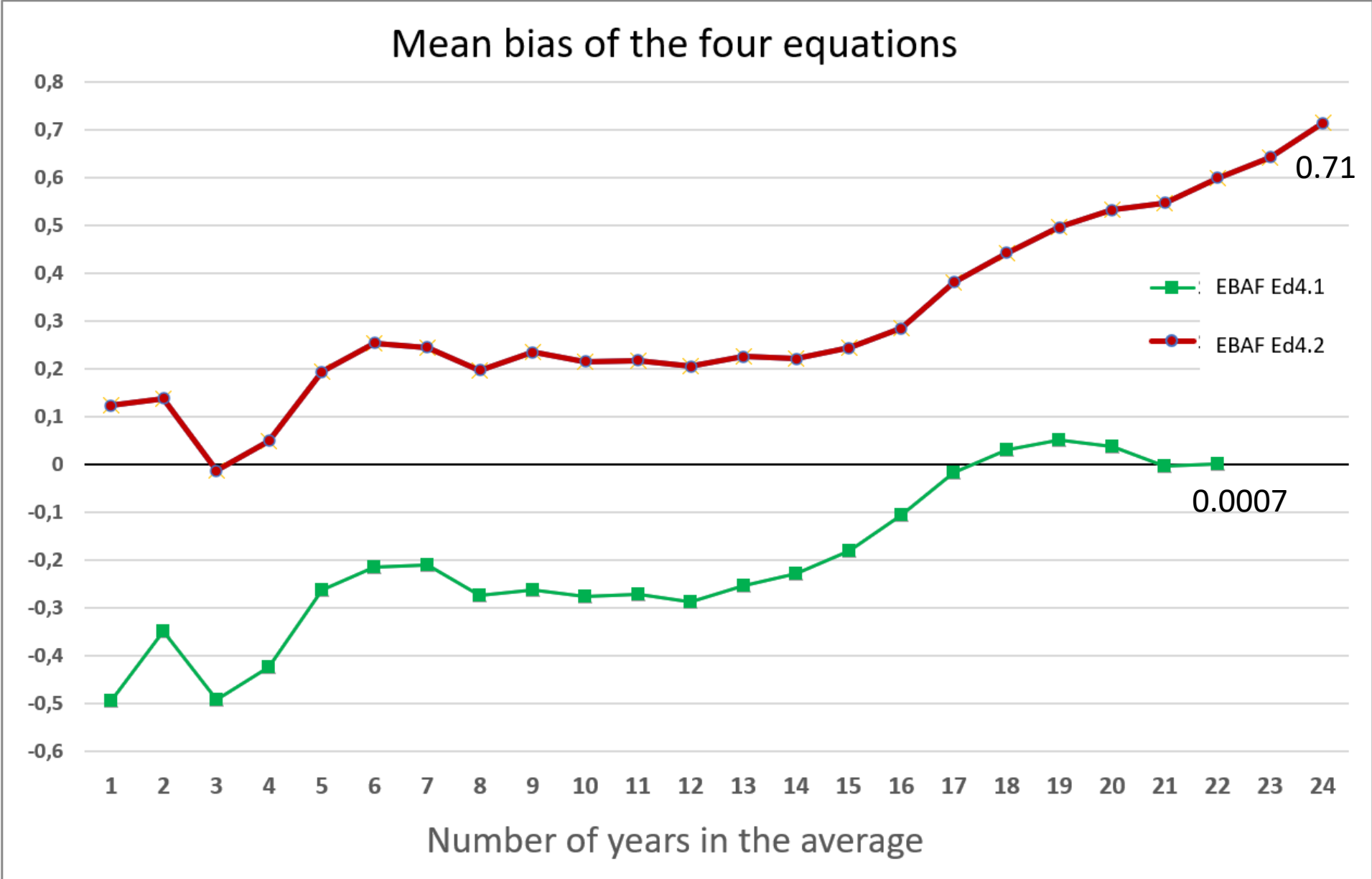
CERES EBAF Ed4.1	Version 3, 22 years	(April 2000 – March 2022) (Wm ⁻²)	Mean:	0.0007
CERES EBAF Ed4.2	Version 4, 22 years	(April 2000 – March 2022) (Wm ⁻²)		0.5994
CERES EBAF Ed4.2	Version 4, 24 years	(April 2000 – March 2024) (Wm ⁻²)		0.7147
CERES EBAF Ed4.2.1	Version 1, 22 years	(April 2000 – March 2022) (Wm ⁻²)		0.5995
CERES EBAF Ed4.2.1	Version 1, 24 years	(April 2000 – March 2024) (Wm ⁻²)		0.7036

Eq. (1)	SFC SW down	– SW up	+ LW down	– LW up	(clear) = TOA LW (clear)/2	
	240.8680	– 29.0724	+ 317.4049	– 398.5211	= 266.0122 /2	– 2.3267
	241.0969	– 29.7521	+ 317.8744	– 398.5890	= 265.9594 /2	– 2.3495
	241.0514	– 29.7043	+ 318.0984	– 398.7742	= 265.9748 /2	– 2.3161
	241.0969	– 29.7521	+ 317.8744	– 398.5890	= 265.9594 /2	– 2.3495
	241.0362	–29.6972	+ 318.1063	– 398.8026	= 265.9732/2	– 2.3439

Eq. (2)	SFC SW down	– SW up	+ LW down	(clear) = 2 × TOA LW (clear)	
	240.8680	– 29.0724	+ 317.4049	= 2 × 266.0122	– 2.8238
	241.0969	– 29.7521	+ 317.8744	= 2 × 265.9594	– 2.6996
	241.0514	– 29.7043	+ 318.0984	= 2 × 265.9748	– 2.5042
	241.0969	– 29.7521	+ 317.8744	= 2 × 265.9594	– 2.6996
	241.0362	–29.6972	+ 318.1063	=2×265.9732	– 2.5012

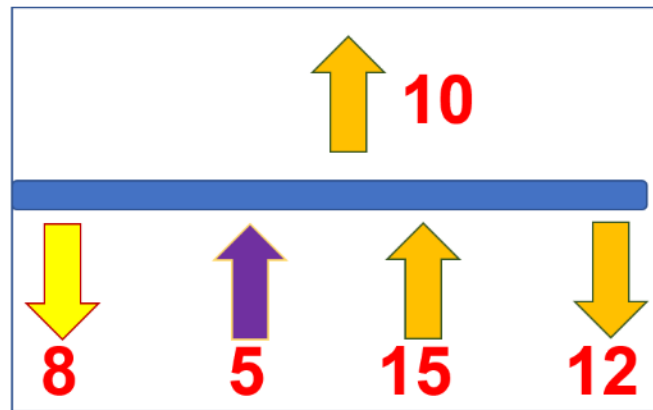
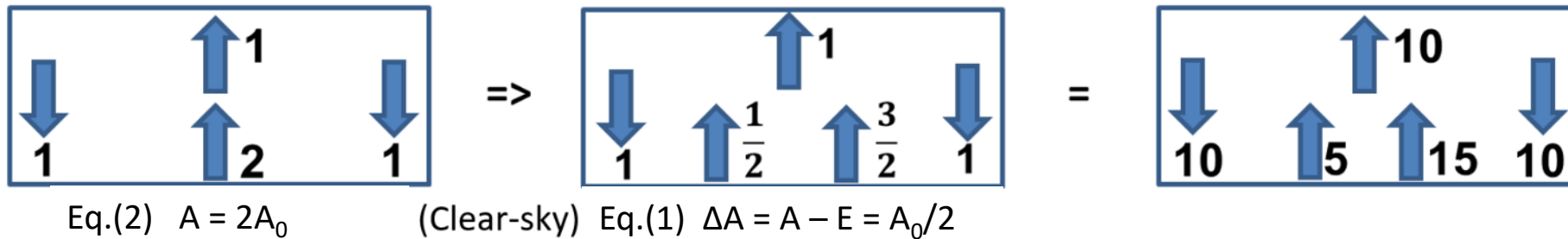
Eq. (3)	SFC SW down	– SW up	+ LW down	– LW up	(all) = [TOA LW (all) – LWCRE]/2	
	186.8544	– 23.1629	+ 345.0108	– 398.7550	= (240.2450 – 25.7672)/2	+ 2.7083
	187.1451	– 23.4950	+ 346.1057	– 398.4220	= (240.3317 – 25.6277)/2	+ 3.9818
	187.1756	– 23.4607	+ 346.3158	– 398.6162	= (240.3894 – 25.5854)/2	+ 4.0126
	187.1451	– 23.4950	+ 346.1059	– 398.4220	= (240.3317 – 25.6277)/2	+ 3.9818
	187.1513	– 23.4547	+ 346.3226	– 398.6131	= (240.3894 – 25.5835)/2	+ 4.0032

Eq. (4)	SFC SW down	– SW up	+ LW down	(all) = 2 × TOA LW (all) + LWCRE	
	186.8544	– 23.1629	+ 345.0108	= 2 × 240.2450 + 25.7672	+ 2.4450
	187.1451	– 23.4950	+ 346.1057	= 2 × 240.3317 + 25.6277	+ 3.4647
	187.1756	– 23.4607	+ 346.3158	= 2 × 240.3894 + 25.5854	+ 3.6665
	187.1451	– 23.4950	+ 346.1059	= 2 × 240.3317 + 25.6277	+ 3.4649
	187.1513	– 23.4547	+ 346.3226	= 2 × 240.3894 + 25.5835	+ 3.6565



The **N**-numbers, as solution of the equations

Pure geometry: No reference to GHGs



$$8 + 12 - 15 = 10 / 2$$

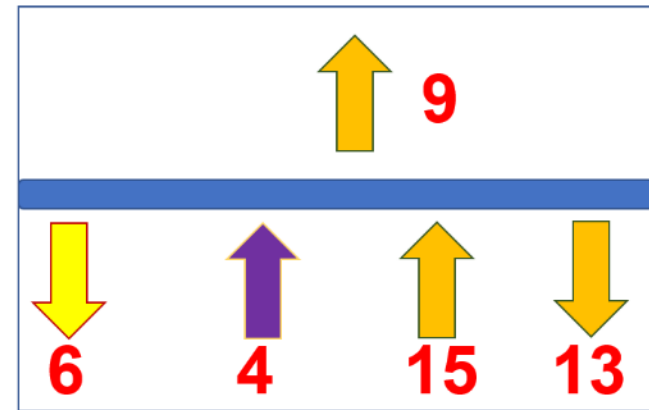
$$8 + 12 = 10 \times 2$$

Eq. (1) SFC Net = $A_0/2$

Eq. (2) SFC Tot = $2A_0$

Clear-sky

$L = 1$
 \Rightarrow



$$6 + 13 - 15 = (9 - 1)/2$$

$$6 + 13 = 9 \times 2 + 1$$

Eq. (3) SFC Net = $(A_0 - L)/2$

Eq. (4) SFC Tot = $2A_0 + L$

All-sky

The flux components with LWCRE = 1

TOA LW	clear-sky = 10	TOA LW	all-sky = 9
SFC LW up	clear-sky = 15	SFC LW up	all-sky = 15
SFC LW down	clear-sky = 12	SFC LW down	all-sky = 13
SFC LW net	clear-sky = -3	SFC LW net	all-sky = -2
SFC SW net	clear-sky = 8	SFC SW net	all-sky = 6
SFC SW+LW net	clear-sky = 5	SFC SW+LW net	all-sky = 4
SFC SW+LW total	clear-sky = 20	SFC SW+LW total	all-sky = 19
G greenhouse effect	clear-sky = 5	G greenhouse effect	all-sky = 6
SWCRE (surface)	= -2	LWCRE (surface, TOA)	= 1

Fit model to observation: CERES EBAF Ed4.2.1, 24 years, Oct 2000 — Sep 2024 data

Best fit: 1 unit = 1 = LWCRE = $26.68 \pm 0.01 \text{ Wm}^{-2}$

CERES EBAF Ed4.2.1, 24 years
 October 2000 – September 2024, **1** = 26.68 Wm⁻²

		N	N × Unit	EBAF Ed4.2.1	Difference (Wm ⁻²)
Clear-Sky TOA	LW	40 /4	266.80	265.98	-0.82
	SW	8 /4	53.36	53.76	0.40
	Net	3 /4	20.01	20.46	0.45
Clear-Sky Surface	LW down	12	320.16	318.25	-1.91
	LW up	15	400.20	398.92	-1.28
	LW net	-3	-79.98	-80.67	0.63
	SW down	9	240.12	240.97	0.85
	SW up	1	26.68	29.67	2.99
	SW net	8	213.44	211.30	-2.14
	SW + LW net	5	133.40	130.63	-2.77
	G	5	133.40	132.94	-0.46

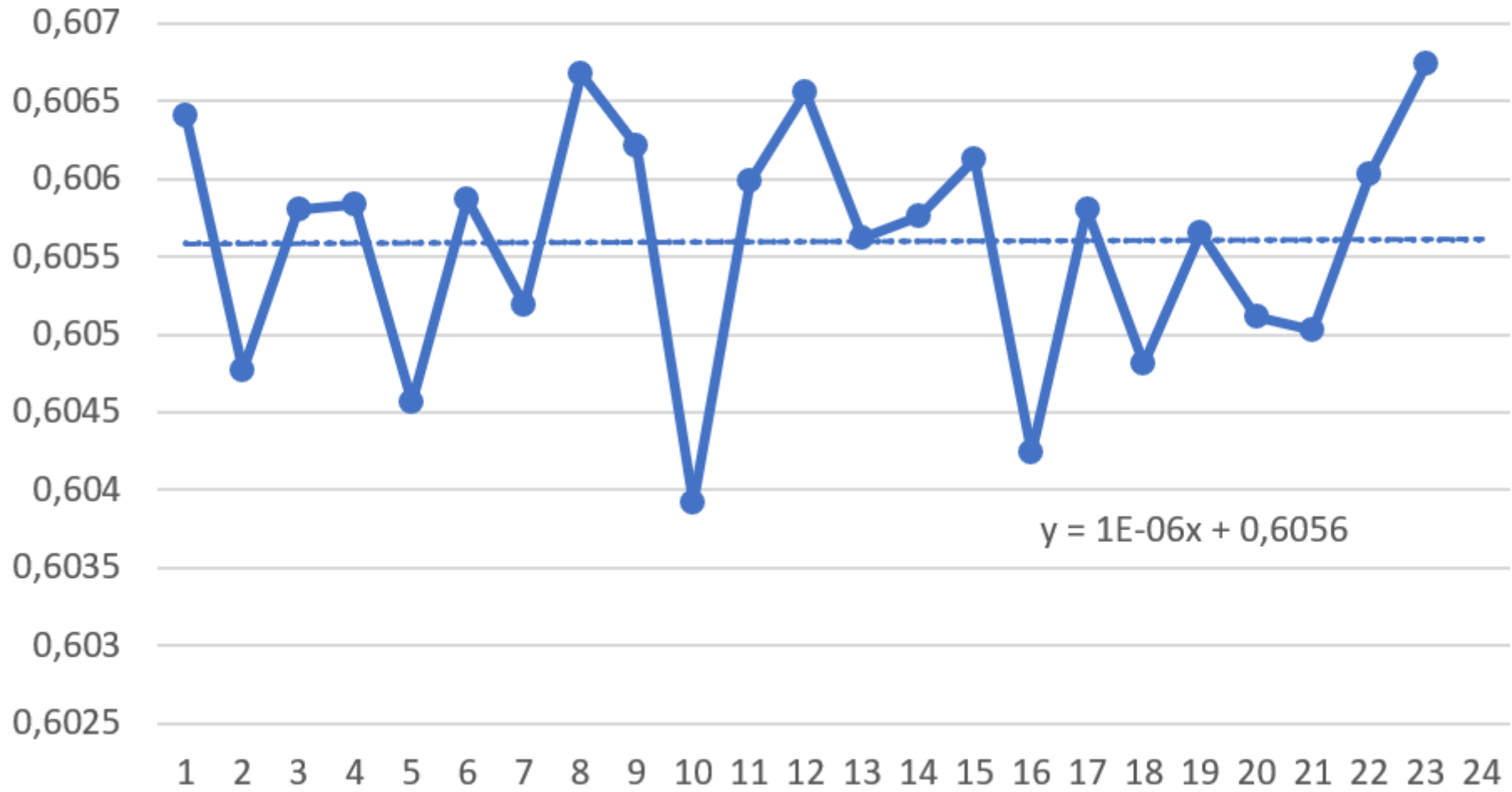
CERES EBAF Ed4.2.1, 24 years; **1** = 26.68 Wm⁻²

	All-sky	N	N × Unit	EBAF Ed4.2.1	Data – Theory (Wm ⁻²)
TOA	SW insolation	51 /4	340.17	340.21	0.04
	SW up	15 /4	100.05	98.91	-1.14
	LW up	36 /4	240.12	240.42	0.30
	TOT net	0	0	0.88	0.88
Surface	SW down	7	186.76	187.12	0.36
	SW up	1	26.68	23.44	-3.24
	SW net	6	160.08	163.68	3.60
	LW down	13	346.84	346.44	-0.40
	LW up	15	400.20	398.72	-1.48
	LW net	-2	-53.36	-52.28	1.08
	TOT net	4	106.72	111.40	4.68
	CRE				
TOA	SW	-7 /4	-46.69	-45.15	1.54
	LW	1	26.68	25.56	-1.12
	TOT	-3 /4	-20.01	-19.59	0.42

g(clear), CERES EBAF Ed4.2.1, 24-yr mean (Oct 2000 – Sep 2024)

Clear-sky	N	Theory	CERES
Surface LW up	15	400.20	398.9233
TOA LW up	10	266.80	265.9823
G	5	133.40	132.9410
g	5/15	1/3	0.33325
All-sky	N	Theory	CERES
Surface LW up	15	400.20	398.72
TOA LW up	9	240.12	240.42
G	6	160.08	158.30
g	6/15	0.4	0.397

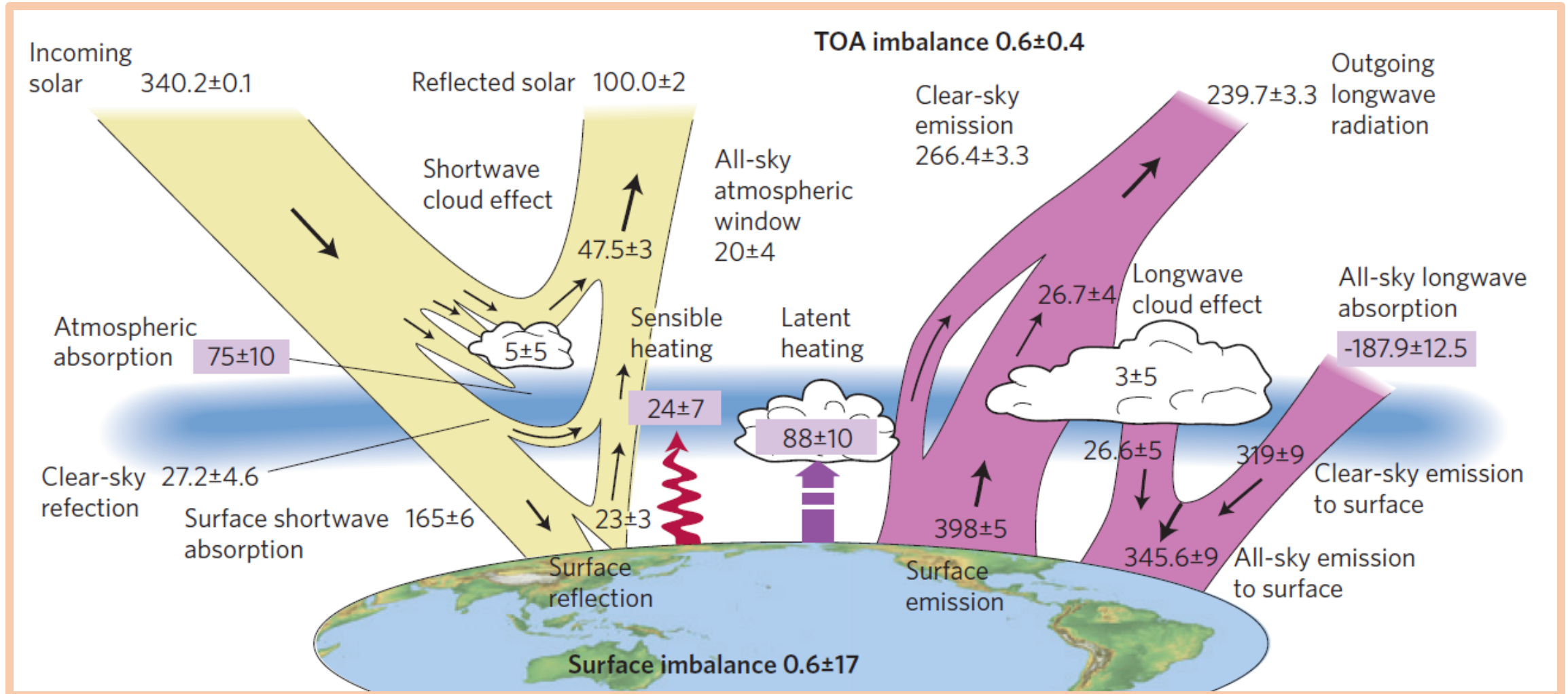
ASR/ULW

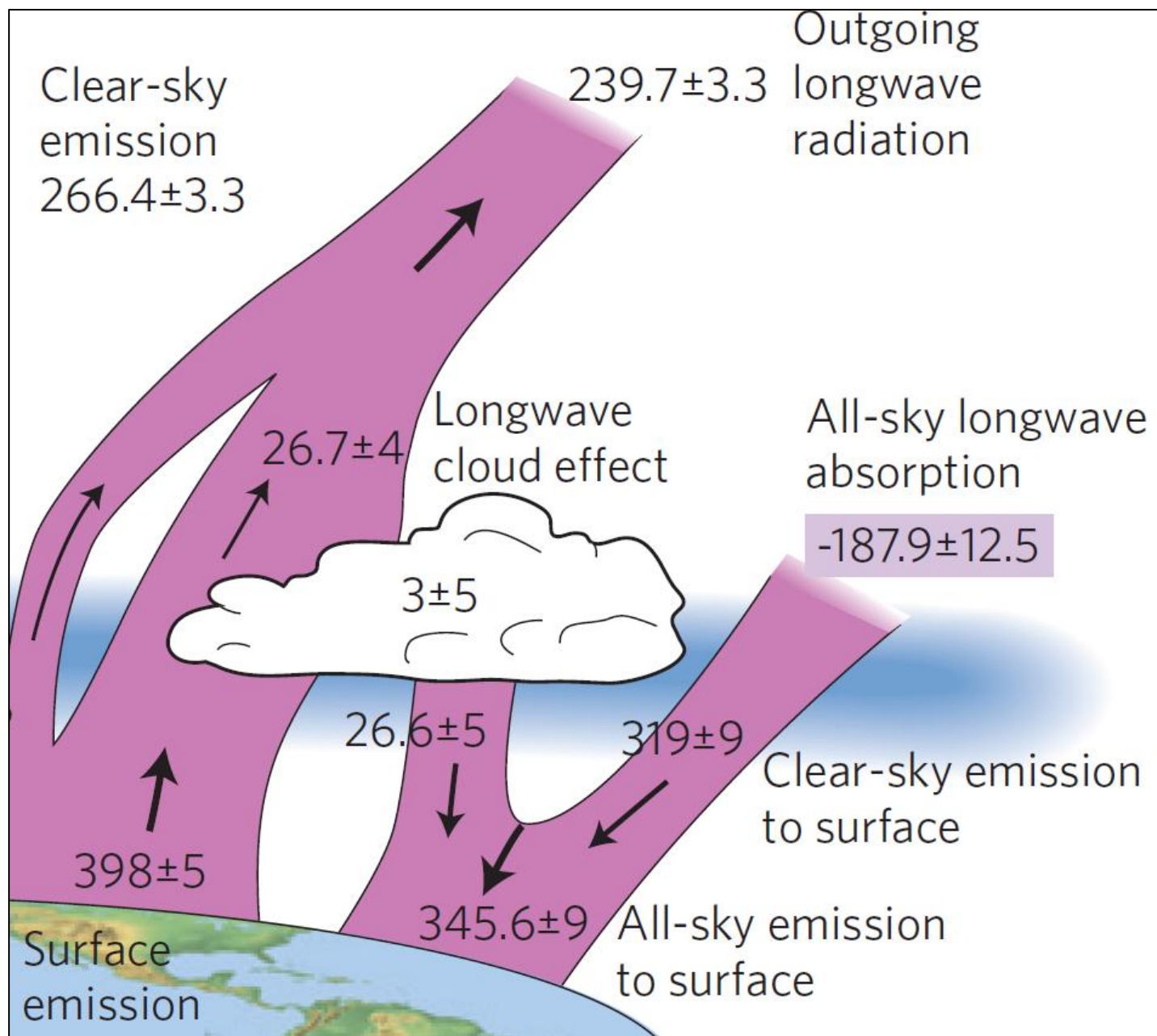


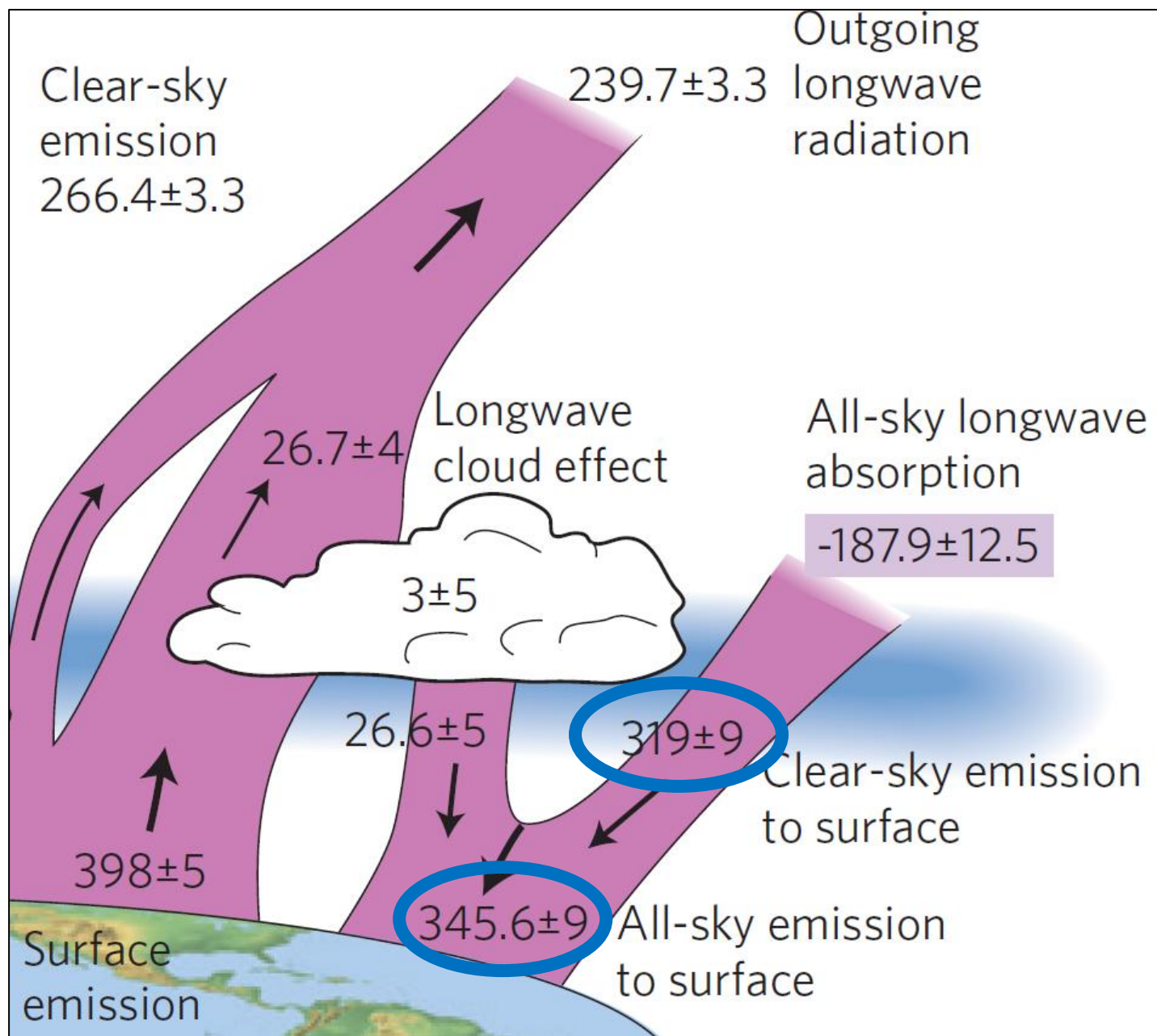
An update on Earth's energy balance in light of the latest global observations

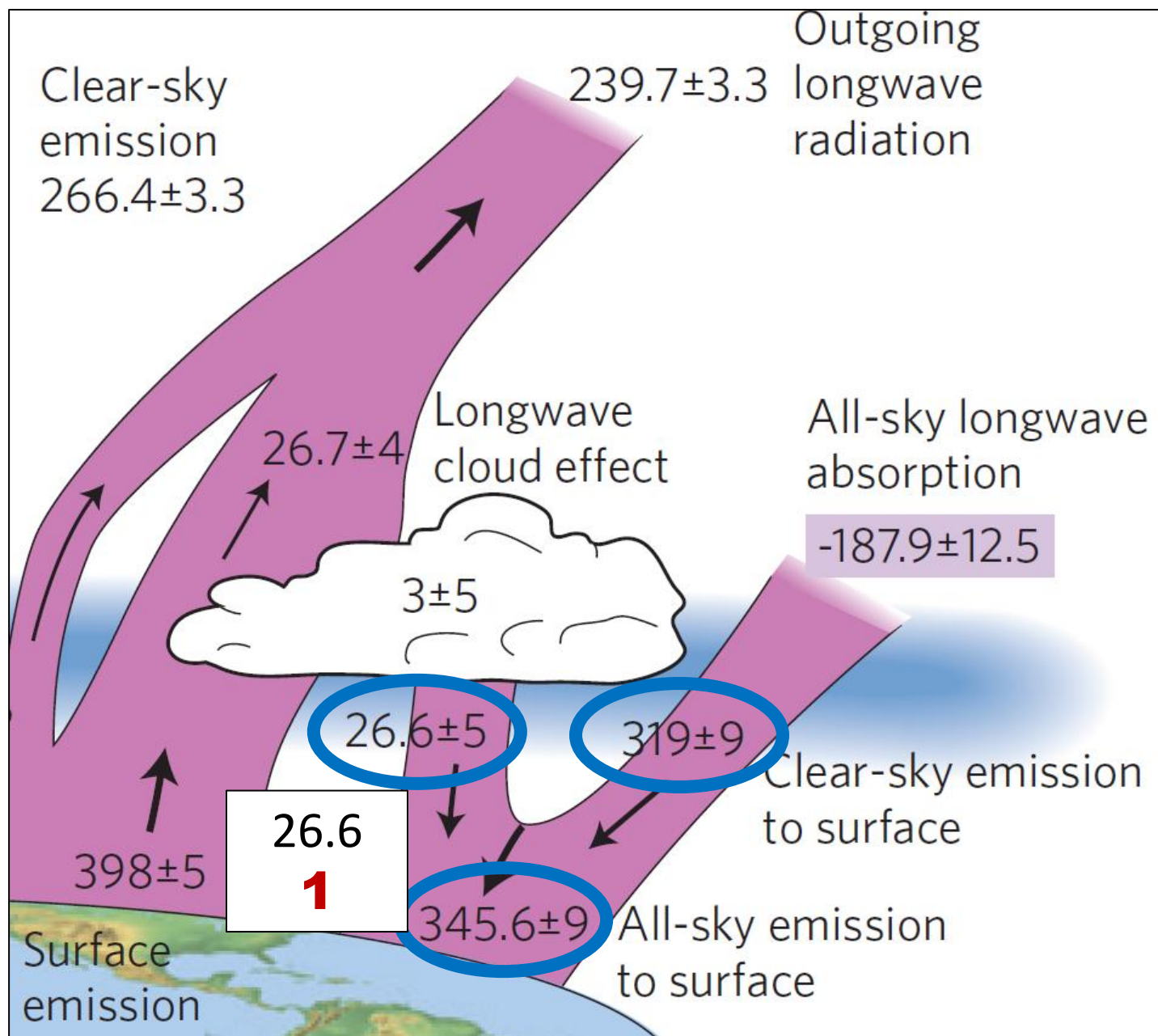
Nature Geosci. (2012)

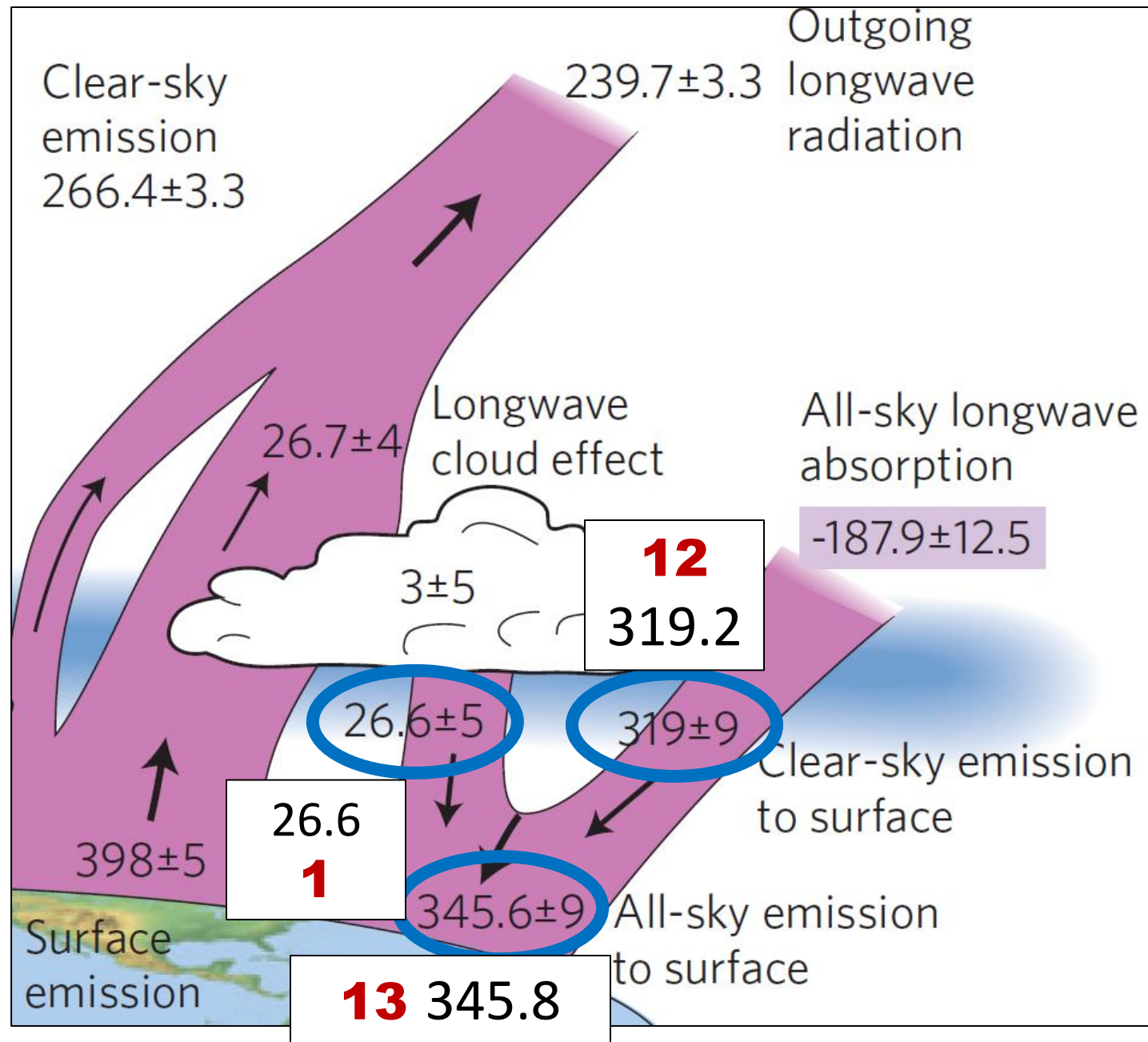
Graeme L. Stephens^{1*}, Juilin Li¹, Martin Wild², Carol Anne Clayson³, Norman Loeb⁴, Seiji Kato⁴, Tristan L'Ecuyer⁵, Paul W. Stackhouse Jr⁴, Matthew Lebsock¹ and Timothy Andrews⁶

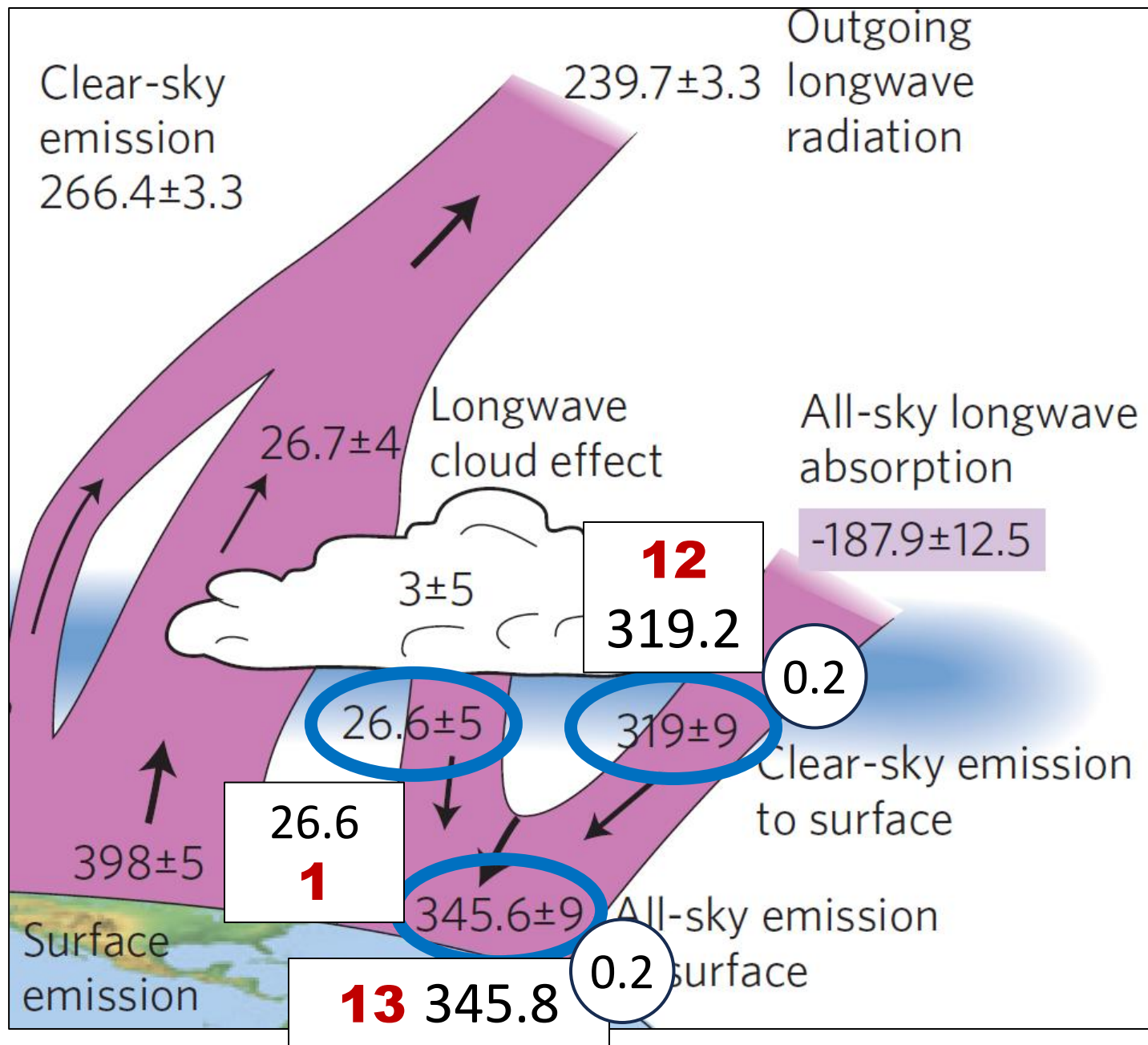


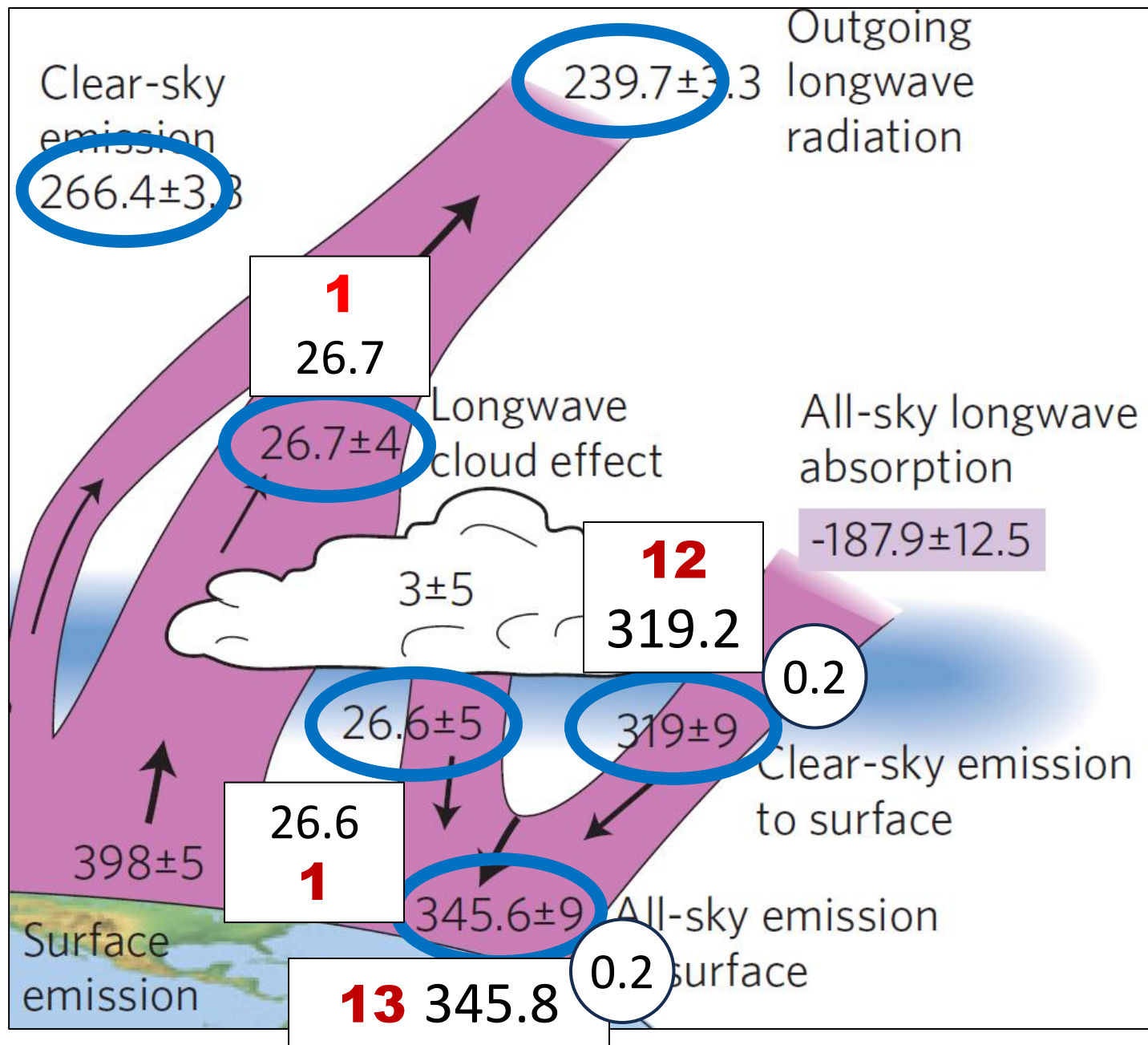


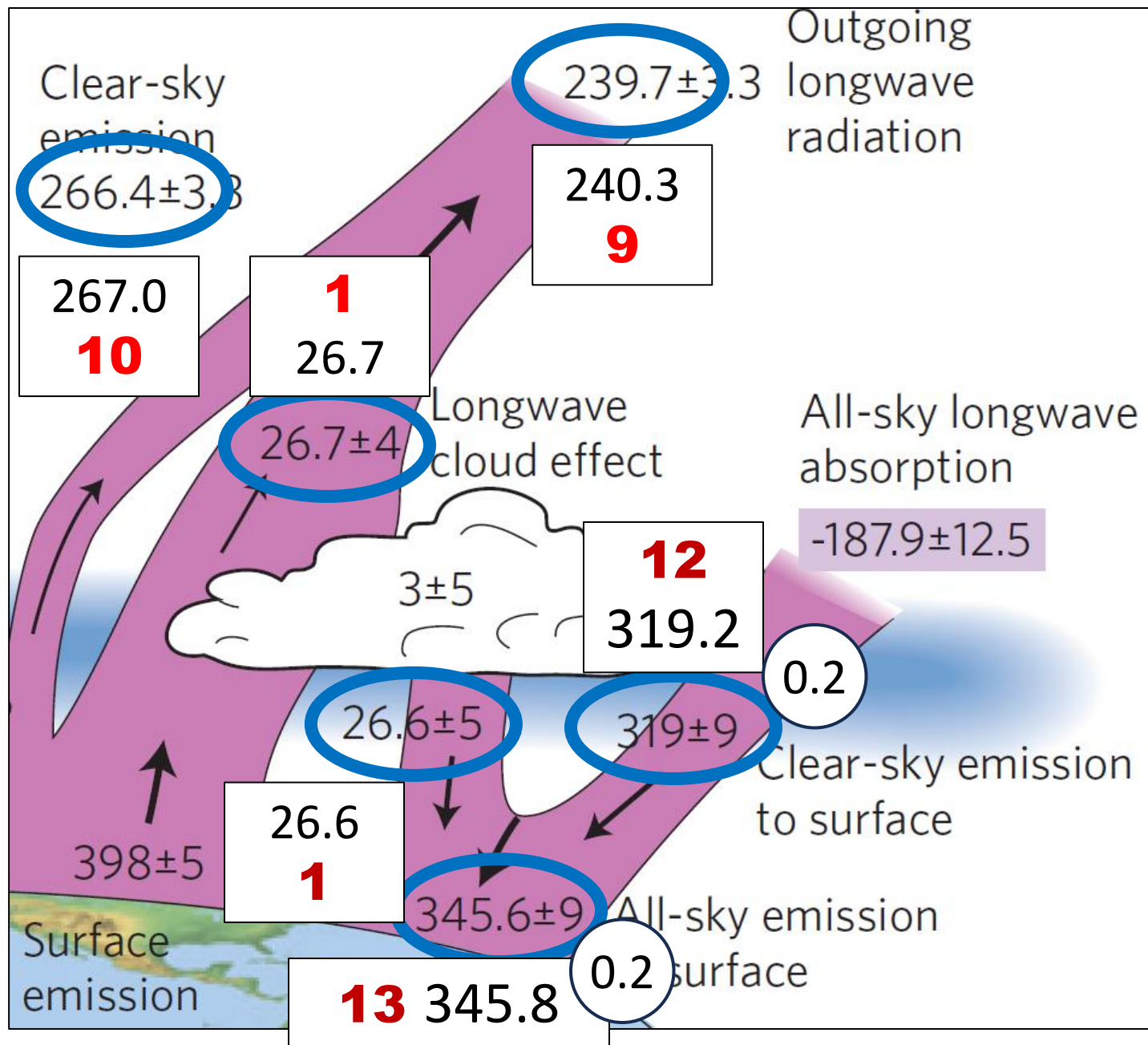


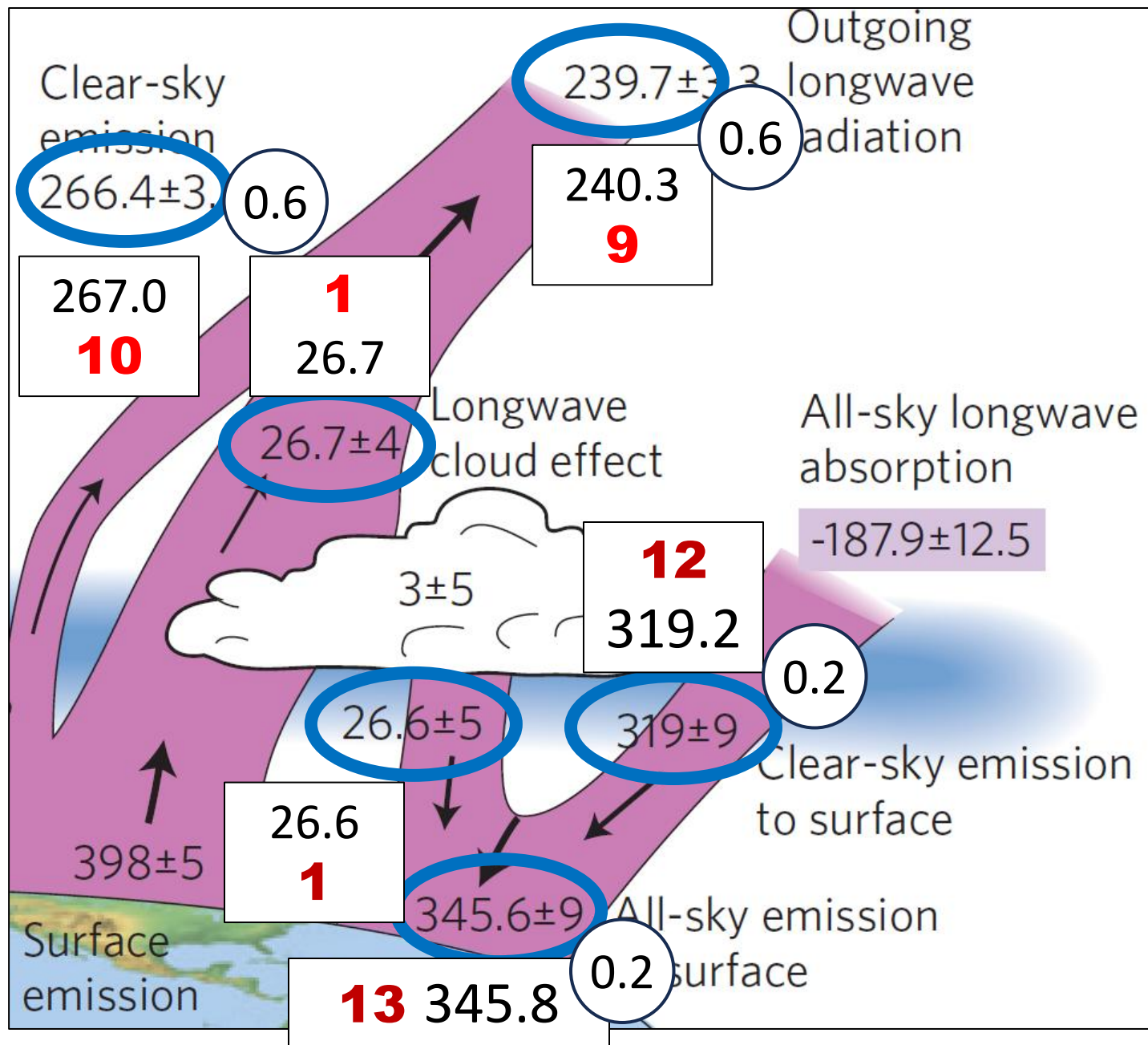


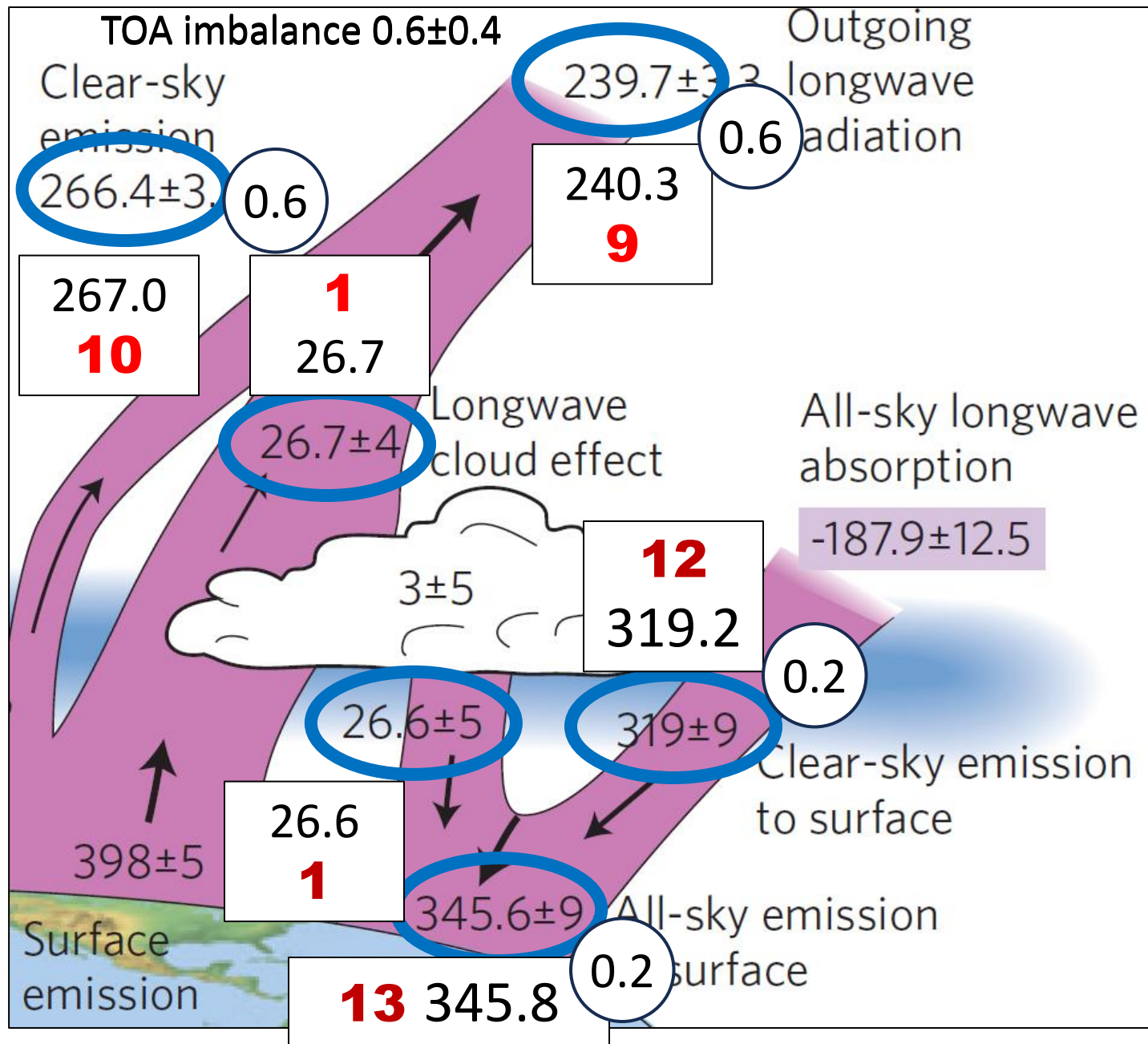


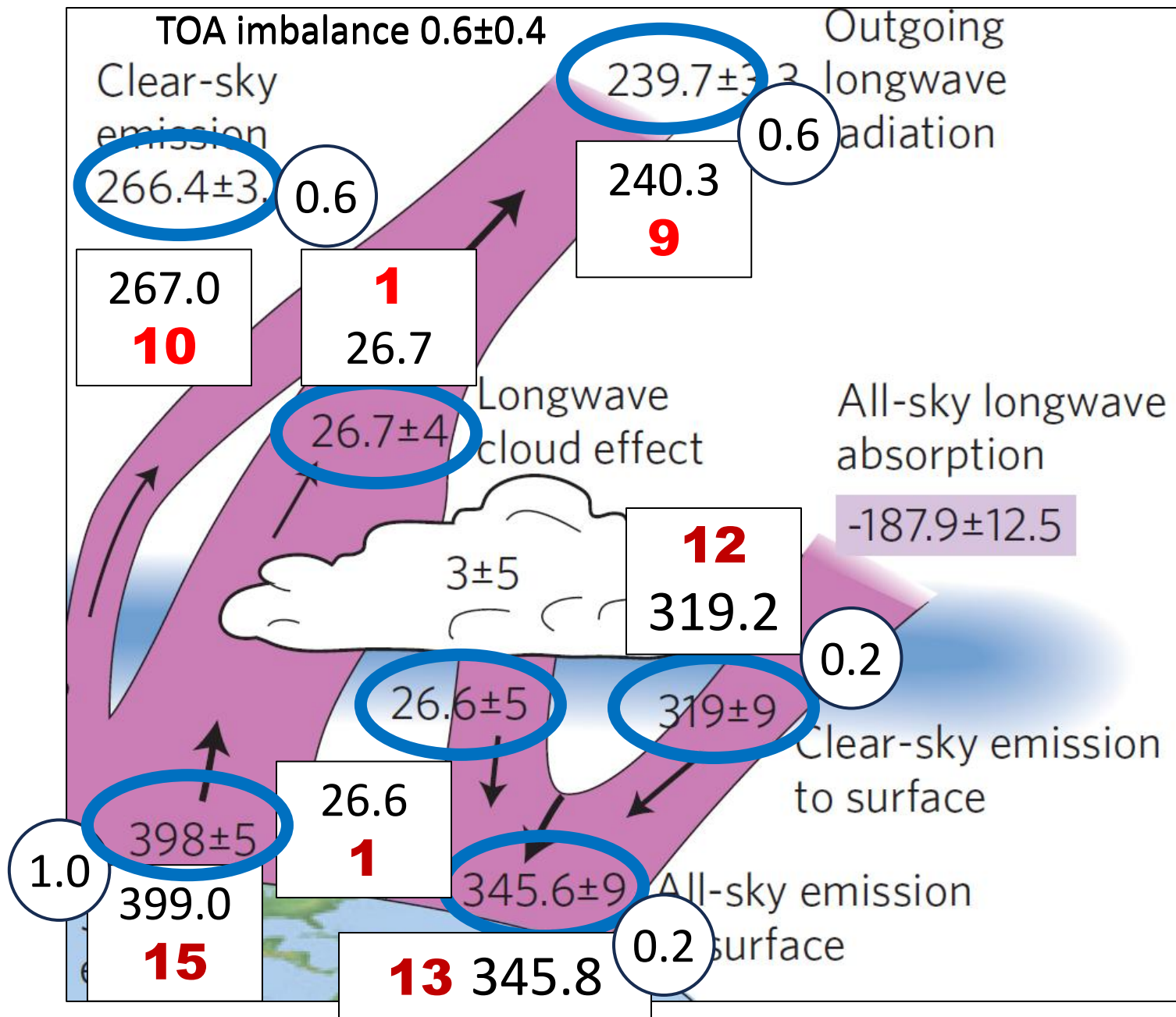












TSI
51

Earth Energy Imbalance (EEI) = 0.54 ± 0.3

LWCRE

1

26.68

TSI

51

1360.68

Incoming Solar

340.2 ± 0.1

51/4

340.17

Reflected Solar

100.2 ± 2.4

15/4

100.05

Outgoing LW

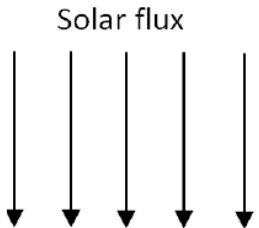
239.5 ± 2.4

36/4

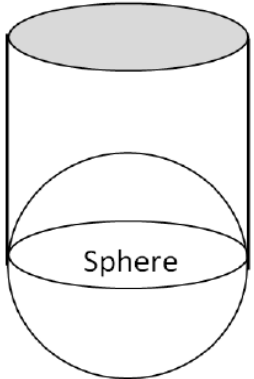
240.12

Eq. (3) Sensible heat + Evaporation = (Outgoing LW - LWCRE)/2
 25.4 + 81.1 = (239.5 - 26.68)/2 + 0.09

Eq. (4) Absorbed SW + All-sky emission = 2 × Outgoing LW + LWCRE
 160.7 + 345.1 = 2 × 239.5 + 26.68 + 0.12



Disk



186.76 **7**

Surface SW

184.0 ± 5.6

26.68 **1**

Sensible heat

25.4 ± 9.4

80.04 **3**

Evaporation

81.1 ± 6.1

Precipitation

81.1 ± 6.1

Absorbed SW

160.7 ± 5.3

160.08 **6**

Surface Reflection

23.3 ± 2.1

26.68 **1**

Surface emission

400.7 ± 4.8

400.20 **15**

346.84 **13**

All-sky emission

345.1 ± 5.7

Stephens et al. (2023, BAMS)

TSI
51

Earth Energy Imbalance (EEI) = 0.54 ± 0.3

Incoming Solar
 340.2 ± 0.1

Reflected Solar
 100.2 ± 2.4

Outgoing LW
 239.5 ± 2.4

LWCRE

1

26.68

26.7 ± 4 Longwave cloud effect

Stephens et al. (2012)

TSI
51

1360.68

0.03

51/4
340.17

0.15

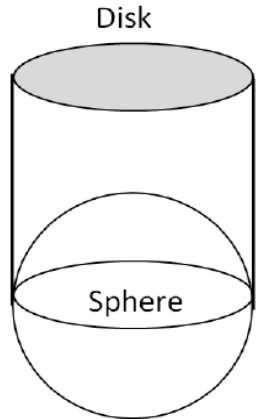
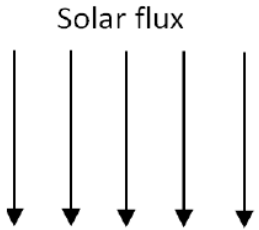
15/4
100.05

-0.62

36/4
240.12

Eq. (3) Sensible heat + Evaporation = (Outgoing LW - LWCRE)/2
 $25.4 + 81.1 = (239.5 - 26.68)/2 + 0.09$

Eq. (4) Absorbed SW + All-sky emission = $2 \times$ Outgoing LW + LWCRE
 $160.7 + 345.1 = 2 \times 239.5 + 26.68 + 0.12$



Precipitation

81.1 ± 6.1

186.76 **7**

2.76 Surface SW
 184.0 ± 5.6

1.28

26.68 **1**
Sensible heat
 25.4 ± 9.4

80.04 **3**

Evaporation
 81.1 ± 6.1

-1.06

Surface Reflection
 23.3 ± 2.1

Surface emission
 400.7 ± 4.8

346.84 **13**

0.62 Absorbed SW
 160.7 ± 5.3
160.08 **6**

26.68 **1**

-3.38

0.50

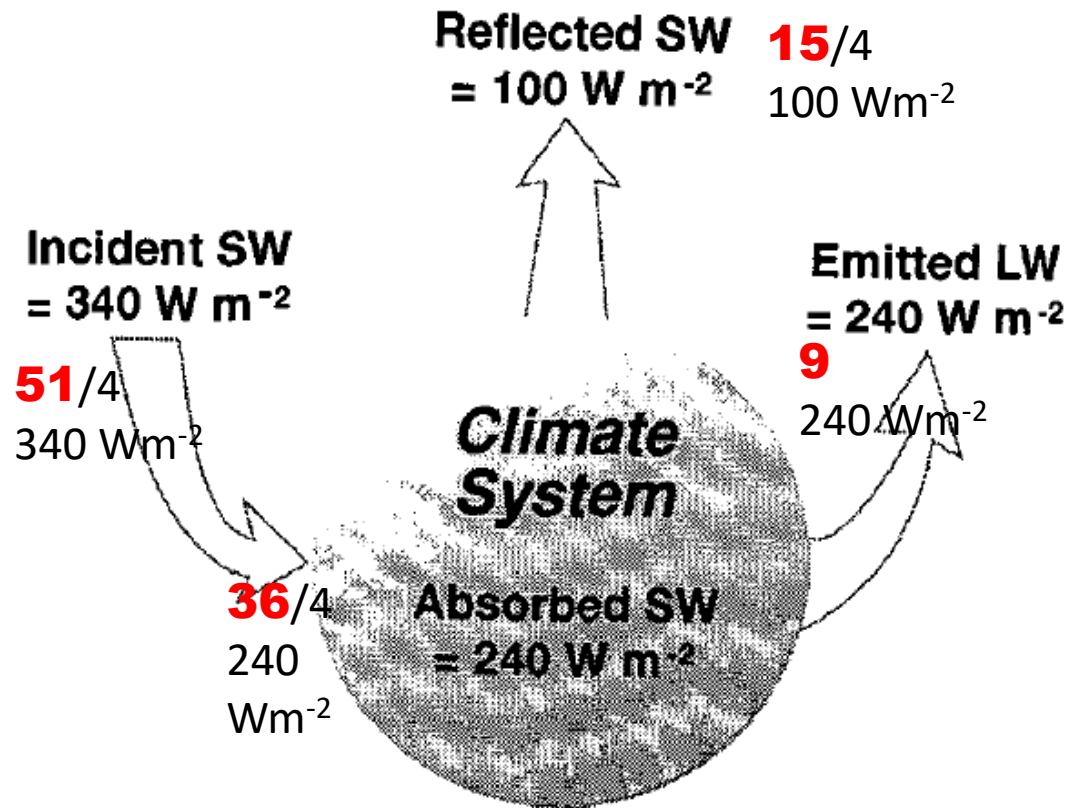
400.20 **15**

1.74

All-sky emission
 345.1 ± 5.7

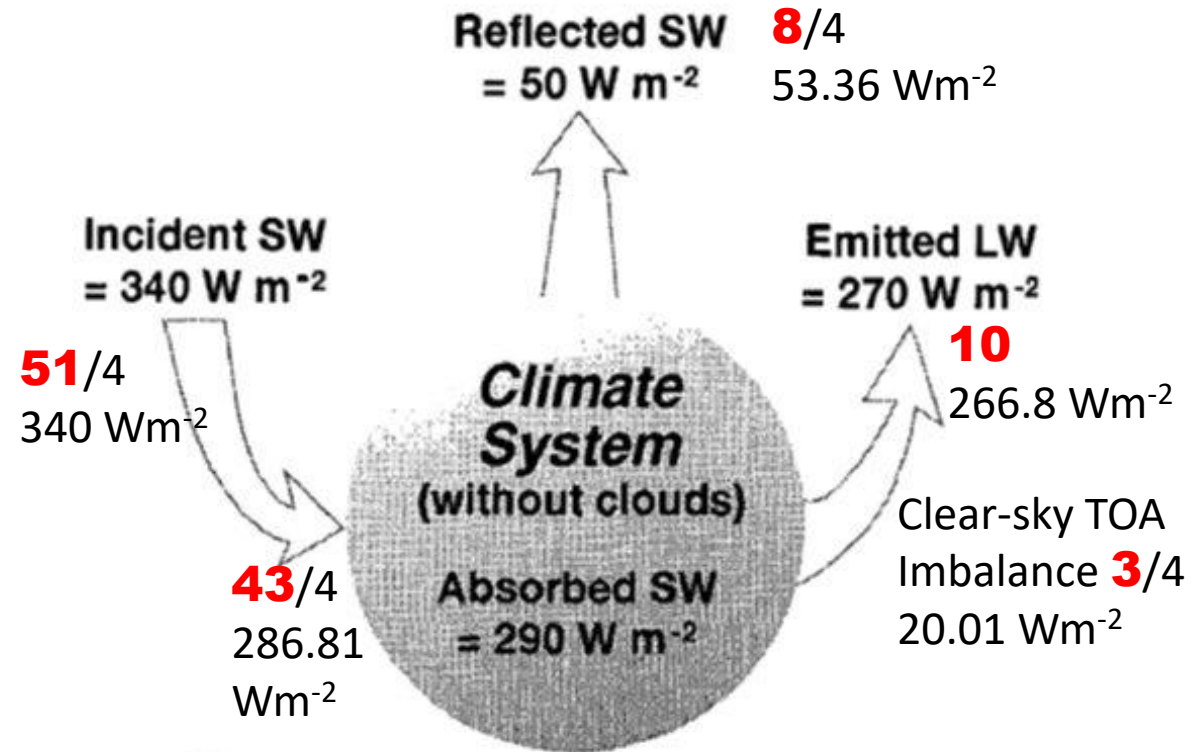


Wielicki et al. (1995)



Theory verified

Incident SW = **51/4** = 340 W m^{-2}
 Reflected SW = **15/4** = 100 W m^{-2}
 Absorbed SW = **36/4** = 240 W m^{-2}
 Emitted LW = **36/4** = 240 W m^{-2}



Cloud-Radiative Forcing (CRF)

LW CRF = 30 W m^{-2}
 SW CRF = -50 W m^{-2}
 NET CRF = -20 W m^{-2}

Theory verified

LW CRF = **1** = 26.68 W m^{-2}
 SW CRF = **-7/4** = -46.69 W m^{-2}
 Net CRF = **-3/4** = -20.01 W m^{-2}

Concluding

- „*CERES is an investigation to examine the role of cloud/radiation feedback in the Earth's climate system.*” (Wielicki et al. 1996). The investigation concludes in a set of four radiative transfer equations with a solution of small integer ratios related to the longwave radiative effect of clouds (LWCRE).
- CERES CDR (climate data record) prove the validity of the four equations and justify their integer solution within uncertainty; inversely, CERES CDR prove that Earth follows accurately these four equations and implements their integer solution system.
- The four equations and the extended set of their integer solution give an explicit answer to the initial problem, by quantifying the SW, LW and net CRE in the climate system and prescribing their position in the integer ratio system related to the total solar irradiance (TSI).
- The four equations and their integer solution define a steady state of Earth's global mean energy flow system. It is a greenhouse gas-independent geometric state, driven by TSI, and is based on a known radiative transfer constraint: the net radiation at the surface is half of the outgoing LW radiation at the top of the atmosphere in the clear-sky. The all-sky system is modulated by LWCRE.
- The discovery of this steady state has consequences pointing far beyond the original scope of the mission, and affects every aspect of the greenhouse theory: sensitivity studies, cloud feedback estimates, climate models, climate change assessments, and climate projections.
- CERES CDR played an indispensable role in verifying the clear-sky system. Future satellite missions with better absolute calibration are needed to monitor the deviation from the steady-state; to understand the impact of man's activities and predict their long-term consequences.